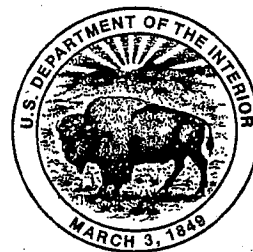


Feasibility Study of Material-Balance Assessment of Petroleum from the New Albany Shale in the Illinois Basin

By M.D. Lewan, J.B. Comer, T. Hamilton-Smith, N.R. Hasenmueller,
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*A material-balance assessment of the petroleum expelled from
source rocks in the New Albany Shale and accountable
petroleum within the Illinois Basin*



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FEASIBILITY STUDY OF MATERIAL-BALANCE ASSESSMENT OF PETROLEUM FROM THE NEW ALBANY SHALE IN THE ILLINOIS BASIN

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ABSTRACT

This study examines the feasibility of assessing petroleum resources of an established petroleum basin by a material-balance method. This method involves determining the amount of petroleum (1) expelled from mature and over-mature source rocks, (2) discovered in place, and (3) lost by migration, surface leakage, and erosion. The Illinois Basin was used for this study because of its extensive drilling history as an established petroleum basin, good estimates of discovered in-place petroleum from production data, well-developed geological framework based on subsurface data, and availability of geochemical databases on source rocks of the New Albany Shale and related petroleum. Although additional research is needed to fully develop this method, the results from this study indicate that a material-balance assessment is feasible. The amount of petroleum expelled from source rocks in the New Albany Shale within the Illinois Basin is estimated to be between 152 and 236 billion barrels. Currently, 12 billion barrels of this expelled petroleum has been discovered in place. Vertical secondary migration along fractures and faults appears to have been dominant in the Illinois Basin. However, 18 to 55 billion barrels of residual petroleum may be lost within carrier beds as a result of some lateral secondary migration. In-place tar deposits of western Kentucky represent 15 to 30 billion barrels of original undegraded petroleum, and an additional 46 to 91 billion barrels of petroleum may have been lost due to erosion of the Upper Pennsylvanian through Permian

section. Although this material balance yields reasonable values, the uncertainty in assumptions and variability in values make it too equivocal at this time to determine the amount of petroleum remaining to be discovered. This feasibility study yields a broad range of -37 to 144 billion barrels of petroleum from the New Albany Shale that remain to be discovered in conventional or unconventional accumulations in the Illinois Basin. The -37 billion barrels indicates that possibly no more undiscovered petroleum remains to be found and that the amounts of accountable petroleum lost to carrier beds or erosion are overstated. The 144 billion barrels indicates that possibly the expulsion factors or the steady-state model for organic richness for the source rocks in the mature part of the basin are overstated. Although this range is broad, it does encompass and does not completely exceed the range of accountable petroleum within the basin. Future research directed toward the uncertainties in petroleum expulsion factors and original organic richness of mature source rocks is attainable and will significantly enhance the utility of this assessment method.

INTRODUCTION

The increasing dependence on foreign petroleum and the decreasing domestic exploration effort in the United States are important national concerns. In order to properly posture domestic and foreign policies to optimize sources of future energy supplies, accurate and credible assessments of petroleum resources in the United States and the world are essential. Assessments of undiscovered petroleum are currently based on statistical evaluations and geological analogs of discovered petroleum. These types of assessments typically reflect amounts of petroleum to be found by existing technologies and exploration strategies and do not necessarily reflect ultimate amounts of unconventional and undiscovered conventional petroleum remaining to be found in a sedimentary basin.

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A more scientifically based method to determine the ultimate petroleum potential of a sedimentary basin may be through a material-balance assessment. This method requires the evaluation of petroleum charge and petroleum losses within a sedimentary basin. Petroleum charge is the amount of petroleum expelled from mature and over-mature source rocks. Petroleum losses are the amounts of petroleum lost as a result of leakage to the surface or dispersed and residual amounts along migration pathways. The difference between these two parameters gives the ultimate petroleum potential of a sedimentary basin. The difference between the ultimate petroleum potential and discovered in-place petroleum gives the amount of undiscovered petroleum remaining within a sedimentary basin. This amount includes both undiscovered conventional and unconventional petroleum.

Determining petroleum charge requires realistic expulsion factors for identified source rocks and a thorough characterization of mature and over-mature source rocks. The first requirement has been problematic in the past because open, anhydrous laboratory pyrolysis methods (i.e., Rock-Eval) are unrepresentative of natural processes and grossly overestimate the amounts of petroleum that are expelled from a source rock (Lewan, 1994). This obstacle has more recently been alleviated in part through the development of a laboratory pyrolysis method (i.e., hydrous pyrolysis) that more closely simulates natural petroleum generation and expulsion (Lewan, 1993). As a result, more realistic expulsion factors for source rocks can be determined. The second requirement, to characterize mature and over-mature source rocks, has not been an obstacle in the past but has often been neglected or oversimplified. Two of the most common transgressions of this requirement are the inclusion of large volumes of organic-lean rocks that have no significant potential as source rocks and the use of only the most organic-rich samples to represent an entire mixed section of source and non-source rock. Currently, knowledge of what constitutes a petroleum source rock and availability of completely cored sections of source rocks have made the fulfillment of this requirement more attainable than in the past.

Determining petroleum losses within a sedimentary basin remains a challenging task. Although considerable efforts have been made to study the expulsion of petroleum from source rocks and the reservoir character of entrapped petroleum, few studies have been made on the intermediate processes responsible for migration of petroleum from source rock to trap. Past attempts to quantify petroleum lost during migration have invoked enormous petroleum losses because of exaggerated ultimate petroleum charge values determined by Rock-Eval pyrolysis. Most of these postulated petroleum losses were attributed to surface leakage, which typically could not be justified by the presence of heavy-oil or tar deposits. In addition, these studies were based on petroleum migration through porous rocks, with only minimal consideration being given to migration along regional fractures or faults. Determining petroleum losses

during migration becomes a more realistic and practical challenge with more realistic values of ultimate petroleum charge, a greater awareness of the consequences of surface leakage, and a new appreciation for migration along fractures and faults.

The overall objective of this study is to determine the feasibility of conducting a material-balance assessment of petroleum within a sedimentary basin. This feasibility is evaluated with respect to amounts and types of data needed, plausibility and utility of assessment values obtained, and timeliness and achievability of future research required. Determining the feasibility of this type of assessment was considered to be best served by studying an established petroleum basin with known quantities of discovered in-place petroleum, sufficient subsurface data for a well-constrained geological framework, and available geochemical data for source-rock characterization. The Illinois Basin (fig. 1) meets these prerequisites with complete assessments of discovered in-place petroleum (Mast and Howard, 1990), a well-established regional stratigraphy (Leighton and others, 1990), and available geochemical databases on the New Albany Shale (Frost, 1980; Robl and others, 1983; Hasenmueller and Leininger, 1987; Comer and others, 1994).

The Illinois Basin has almost 1,700 named fields that produce from approximately 7,000 separate sandstone and carbonate reservoirs (Oltz and others, 1990). The established nature of petroleum production in this basin is reflected in its percentage of stripper production (96 percent), which is more than six times higher than the national average (Oltz and others, 1990). An estimate of the original petroleum in place for known fields in the basin is 12 billion barrels (Mast and Howard, 1990), with an additional 3.4 billion barrels of in-place tar (Noger, 1987; Crysedale and Schenk, 1988). Geochemical correlations and stratigraphic position of reservoirs indicate that more than 99 percent of the produced petroleum was sourced by the New Albany Shale (Hatch and others, 1990). The New Albany Shale consists of varying amounts of brownish-black, organic-rich shale and greenish-gray, organic-lean shale. This formation is present in the subsurface throughout much of the basin, with thicknesses ranging from less than 40 m at the basin margins to more than 140 m in the basin depocenter.

The term petroleum as used in this study refers to naturally occurring liquids that consist predominantly of hydrocarbons (e.g., heavy oil, crude oil, and condensates). Although the definition of petroleum is sometimes expanded to include natural gas (e.g., Levorsen, 1967), natural gas is not considered in this feasibility study, and only the strict definition of petroleum (i.e., from the Latin *petra*, rock and *oleum*, oil) is used. It should also be emphasized that this feasibility study only addresses petroleum expelled from source rocks within the New Albany Shale. Other source rocks within the Illinois Basin have been suggested (Cluff and Byrnes, 1990; Bond and others, 1971), with source rocks within the Scales and Brainard Shales (Ordovician) being



Figure 1. Map of Illinois, Indiana, and western Kentucky showing outline of the Illinois Basin based on outcrop, subcrop, and limit of the New Albany Shale (modified from Cluff and others, 1981). Map also identifies counties referred to in text and wells used in cross sections A-A' and B-B' shown in figure 3.

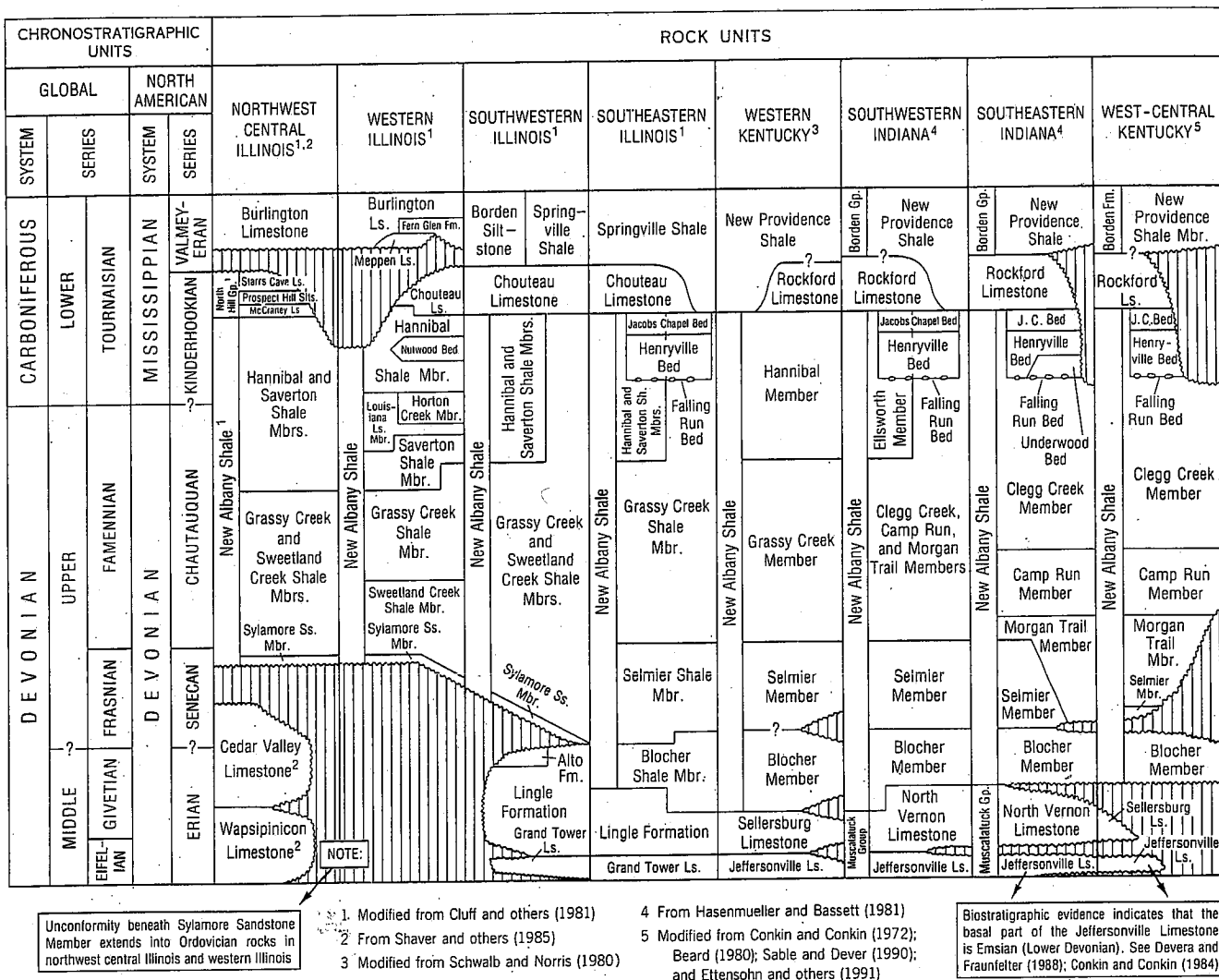


Figure 2. Correlation of New Albany Shale in the Illinois Basin (modified from Hasenmueller and others, 1994). Rock-unit thicknesses are not to scale. Fm., Formation; Gp., Group; J.C., Jacobs Chapel; Ls., Limestone; Mbr., Member; Sh., Shale; Sits., Siltstone; Ss., Sandstone.

prime candidates as additional sources for petroleum (Guthrie and Pratt, 1994) that are not considered in this study.

METHODS AND RESULTS

SOURCE-ROCK CHARACTERIZATION

A petroleum source rock is defined in this study as a rock containing oil-prone organic matter and a total organic carbon (TOC) content equal to or greater than 2.5 weight percent. This TOC limit is based on the petrographic study of Devonian and Mississippian black shales by Lewan (1987). This study showed that a critical quantity of organic matter is needed to generate sufficient bitumen to form a continuous organic network within the matrix of a source rock for expulsion of petroleum to occur. Rocks with oil-prone organic

matter that is dispersed within a mineral matrix may not generate sufficient bitumen to form this continuous network if its original TOC is less than 2.5 weight percent. As a result, these organic-lean rocks may generate but not expel petroleum. Although this TOC limit of 2.5 weight percent is high compared to earlier proposed limits of 0.4 to 1.4 weight percent (Ronov, 1958), higher TOC limits in excess of 2.5 weight percent are now being suggested (Bordenave and others, 1993).

STRATIGRAPHY

The New Albany Shale is a time-transgressive unit extending from Middle Devonian to Early Mississippian (Cluff and others, 1981). Figure 2 gives stratigraphic ages, nomenclature, and correlation of rock units within and

adjacent to the New Albany Shale in different parts of the Illinois Basin. The New Albany Shale overlies Middle Devonian limestones (Sellersburg Limestone, Lingle Formation, and North Vernon Limestone) throughout most of the basin. This basal contact is generally conformable in southern and eastern Illinois, but, farther to the north and west, the basal shale grades laterally into argillaceous limestone, dolostone, and shale of the Lingle and Alto Formations (North, 1969). The New Albany Shale is overlain by Early to Late Mississippian strata throughout most of the basin. However, in northern Illinois, the formation has been partly eroded and is overlain by Pennsylvanian strata. In a small area in extreme southern Illinois, the New Albany Shale is overlain by Upper Cretaceous rocks (Collinson and Atherton, 1975). As shown in figure 2, the New Albany Shale has been locally subdivided into as many as 13 lithologic members.⁶ The regional relationships of these members are shown in the two cross sections of figure 3. Only the members containing organic-rich, brownish-gray to black shales capable of generating petroleum are considered in this study. These members include the Blocher, Selmier, Sweetland Creek, Grassy Creek, Morgan Trail, Camp Run, and Clegg Creek Members. The Blocher, Grassy Creek, Morgan Trail, and Clegg Creek Members consist almost entirely of organic-rich, brownish-black to grayish-black, laminated, pyritic shale. The Selmier, Camp Run, and Sweetland Creek Members consist of varying amounts of organic-lean, greenish-gray shale and organic-rich, brownish-black shale. For a detailed discussion on the stratigraphy of the New Albany Shale, see Hasenmueller and others (1994).

THICKNESS

Thickness of mature and over-mature source rock within a sedimentary basin is a critical parameter in determining the volume of source rock for calculating petroleum charge. An isopach map of the New Albany Shale has been constructed for the Illinois Basin (Cluff and others, 1981), which has been useful in reconstructing paleoenvironments and defining depocenters. Unfortunately, a map of this type includes thicknesses of members that do not contain petroleum source rocks (i.e., Hannibal, Saverton, and Ellsworth Members) and portions of members that are not petroleum source rocks (i.e., Selmier and Camp Run Members). Inclusion of these non-source-rock thicknesses can significantly exaggerate the petroleum charge, and, therefore, an isopach map of only source rocks within the New Albany Shale is

needed. A map of this type currently does not exist in the published literature, but isopach maps of the Blocher Member (fig. 4), Selmier Member (fig. 5), and a composite of the Clegg Creek, Camp Run, Morgan Trail, Grassy Creek, and Sweetland Creek Members (fig. 6) over the entire Illinois Basin are available (Lineback, 1981). The latter composite of members is hereafter referred to informally as the "upper shale composite." Although the use of these maps excludes members consisting of non-source rocks, it does not exclude the thickness of non-source rocks that occur in the Selmier, Camp Run, and Sweetland Creek Members. Corrections to alleviate their inclusion are discussed later under the heading, Organic Richness.

THERMAL MATURITY

The levels of thermal stress experienced by the New Albany Shale within the Illinois Basin have been related to changes in the reflectivity of vitrinite dispersed within organic-rich shales (Barrows and others, 1979, 1980; Barrows and Cluff, 1984; Hamilton-Smith, 1993). Comer and others (1994) have compiled these vitrinite reflectance data and constructed a contour map of the mean random reflectance values. A generalized version of this map is shown in figure 7. Although vitrinite reflectance is a good indicator of thermal stress, it is not always a good indicator of petroleum generation because significant differences in the kinetics of petroleum generation are observed for different types of organic matter (Lewan, 1985; Hunt and others, 1991). This uncertainty is further complicated by the suppressed thermal response of vitrinite reflectance in the New Albany Shale (Nuccio and Hatch, 1994), which also occurs in other Devonian-Mississippian black shales (Lewan, 1985; Price and Barker, 1985). These difficulties were circumvented in this study by establishing a relationship between vitrinite reflectance and the Rock-Eval hydrogen index (mg S₂ hydrocarbons/g total organic carbon) for source rocks in the New Albany Shale at different levels of thermal maturity. The relationship is shown in figure 8 and is based on geochemical data of 61 core samples from 10 wells (table 1). This relationship provides a means of equating the available vitrinite-reflectance contour map (fig. 7) to Rock-Eval hydrogen indices, which are more indicative of petroleum generation. Figure 8 shows that petroleum generation based on the decrease in Rock-Eval hydrogen index occurs between suppressed vitrinite reflectance values of 0.6 and 1.2 percent R_o.

VOLUME

The volume of each source-rock-bearing member isopached by Lineback (1981) was determined for 0.1-percent R_o increments from 0.6 to 1.50 percent R_o. These calculations were made electronically by superimposing the

⁶The Illinois State Geological Survey uses lithic modifiers for the shale members of the New Albany Shale (Blocher Shale Member, Selmier Shale Member, Sweetland Creek Shale Member, Grassy Creek Shale Member, Saverton Shale Member, and Hannibal Shale Member). The Indiana Geological Survey and Kentucky Geological Survey do not use lithic modifiers. For simplicity, this report does not use lithic modifiers.

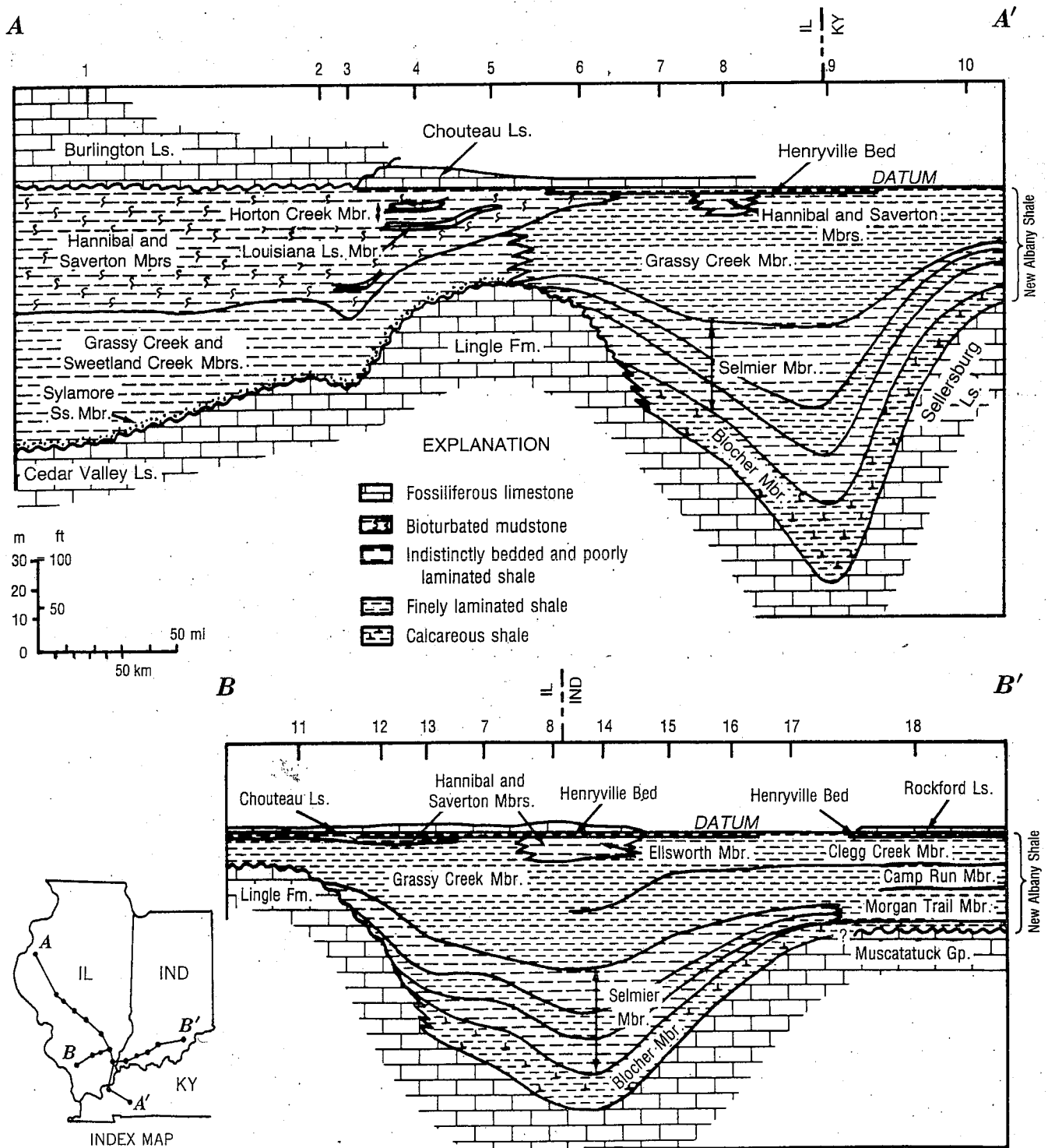


Figure 3. Northwest-southeast (A-A') and west-east (B-B') stratigraphic cross sections of the New Albany Shale showing regional relations among members (modified from Cluff and others, 1981, and Devera and Hasenmueller, 1990). Fm., Formation; Gp., Group; J.C., Jacobs Chapel; Ls., Limestone; Mbr., Member; Slts., Siltstone; Ss., Sandstone. Numbers indicate wells along line of section.

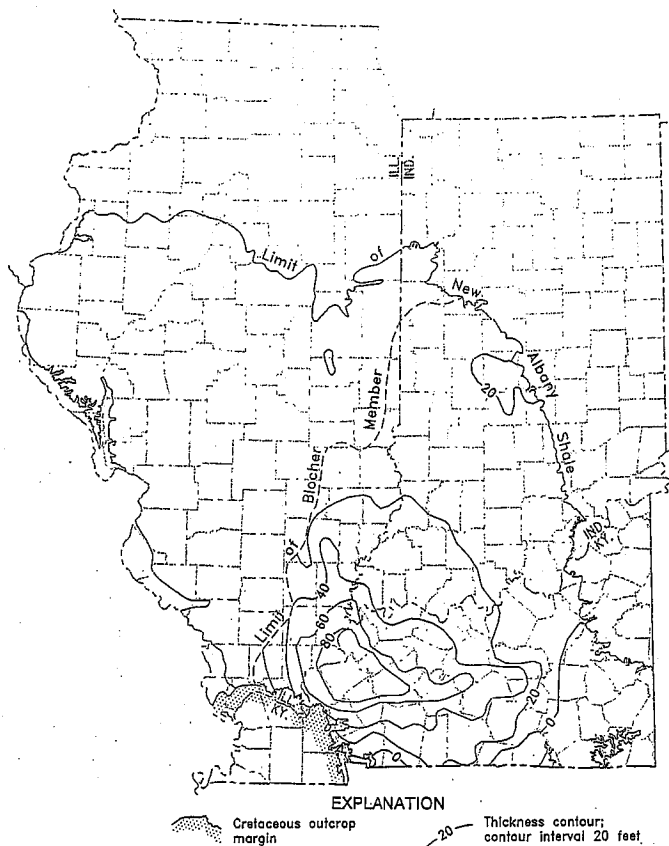


Figure 4. Isopach map showing thickness of Blocher Member of the New Albany Shale in the Illinois Basin (Lineback, 1981).

digitized isorefectance map (fig. 7) over each of the digitized isopach maps of the members of the New Albany Shale (figs. 4, 5, and 6). The isorefectance map was drafted with AutoCAD software (release 12). The isopach maps of members (figs. 4, 5, and 6) were photocopied and then scanned and vectorized using CADcore Tracer software. The digital files of these maps were converted to digital exchange format (DXF) interchange files, using AutoCAD12, and then exported into PC ARC/INFO 3.4D software. In ARC/INFO, all truncated contours were manually joined so that each map consisted of sets of enclosed areas having consistent R_o and thickness values. These areas were assigned attributes that included the lowest R_o value of the two enclosing contours on the R_o map and the mean thickness value of the two enclosing contours on the thickness maps. The thickness maps were then electronically superimposed on the R_o map, causing ARC/INFO to automatically assign combined attributes of area, minimum R_o , and mean thickness to the polygons created by the intersecting R_o and thickness contours. The resulting ARC/INFO attribute table was imported into the relational database, Paradox 3.5, in order to format the final report. In Paradox 3.5, the area of each polygon was multiplied by the mean thickness to obtain the volume. The attributes of R_o , area, and volume were subsequently grouped in the report writer, and the areas and volumes of

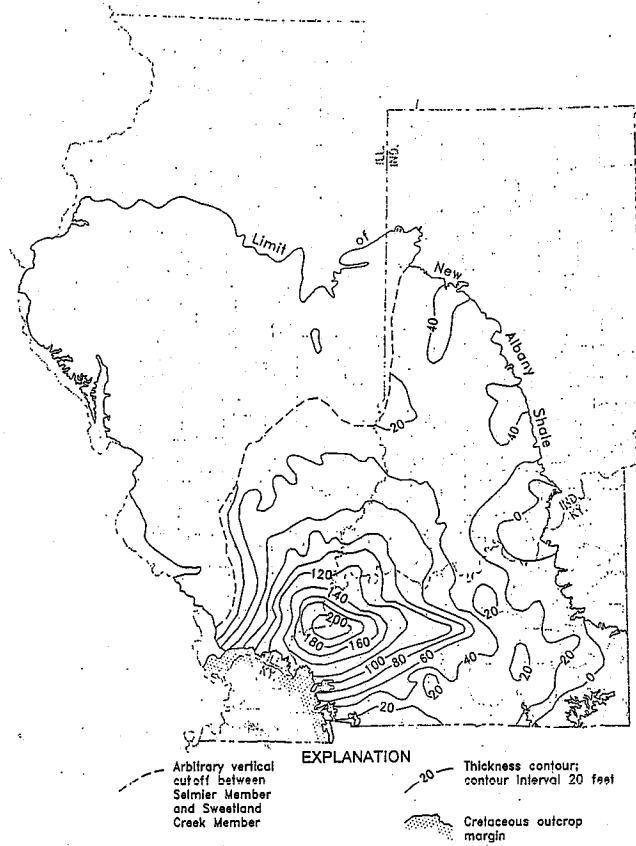


Figure 5. Isopach map showing thickness of Selmier Member of the New Albany Shale in the Illinois Basin (Lineback, 1981). The arbitrary vertical cutoff does not reflect the distribution of the Selmier in northernmost Indiana as recognized by the Indiana Geological Survey.

polygons having the same R_o values were automatically summed and then printed as a report. As a check on this method, area and volume estimates based on the total thickness of New Albany Shale were hand-calculated for Kentucky following the method described in Craft and Hawkins (1959, p. 27-29). The hand-calculated values were found to agree within 10 percent of those calculated using the digital method described above.

The surface areas and total volumes calculated by this method for each member are given in table 2. The total member volumes cannot be directly equated to source-rock volumes because of the non-source-rock intervals that are included in the total thickness of each member. Estimates of source-rock percentages for each member (table 3) were based on total organic carbon data from available cores listed in table 4. Source-rock thicknesses were determined by summing the thickness of rocks containing at least 2.5 weight percent total organic carbon within each member represented in the core. The resulting source-rock thickness was divided by the total thickness of the member represented in the core and multiplied by 100 to obtain a source-rock percentage.

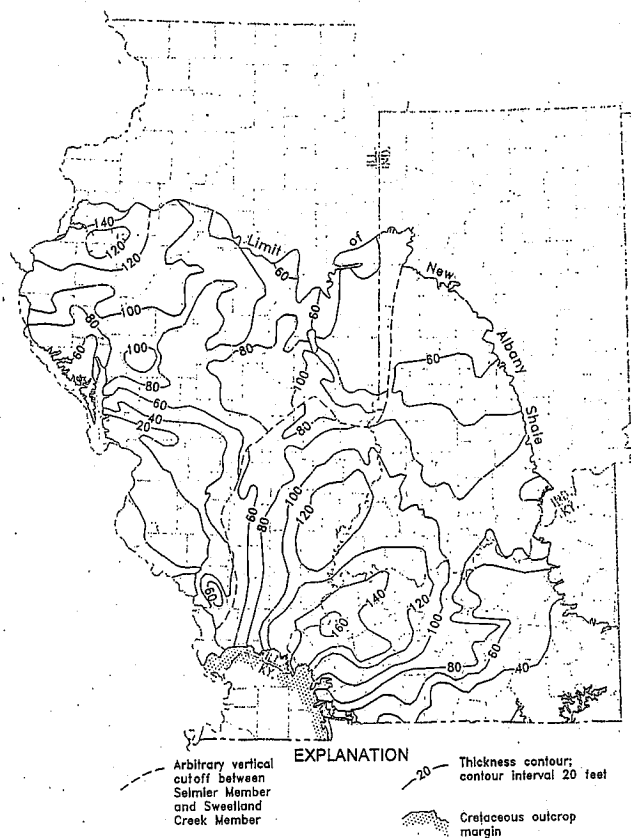


Figure 6. Isopach map showing thickness of the "upper shale composite," which consists of the thickness of the Sweetland Creek and Grassy Creek Members of the New Albany Shale west of the limit of the Selmier Member, the thickness of the Grassy Creek Member only where the Selmier Member is recognized in Illinois and Kentucky, and the combined thickness of the Morgan Trail, Camp Run, and Clegg Creek Members of the New Albany Shale in Indiana (Lineback, 1981). The arbitrary vertical cutoff does not reflect the distribution of the Selmier in northernmost Indiana as recognized by the Indiana Geological Survey.

As shown in table 3, the composite of upper shale members has a narrow range of source-rock percentages from 92 to 100 percent, with no systematic change between vitrinite reflectances of 0.44 and 0.95 percent R_o . The mean source-rock percentage of 97.5 percent obtained from these cores was used for all of the upper shale composite to determine volume of source rock in this mapped unit as shown in table 2. The total calculated volume of source rock in this upper shale composite is 1,119 km^3 , with more than 60 percent of this volume occurring in the 0.6- to 0.7-percent R_o interval (table 2). Inclusion of the non-source rocks of the Sweetland Creek Member in this upper shale composite does not influence these source-rock volumes because this member only occurs at vitrinite reflectance levels of less than 0.6 percent R_o . The Blocher Member also has a narrow range of source-rock percentages from 98 to 100 percent from vitrinite reflectances of 0.44 to 0.70 percent R_o (table 3). A mean of

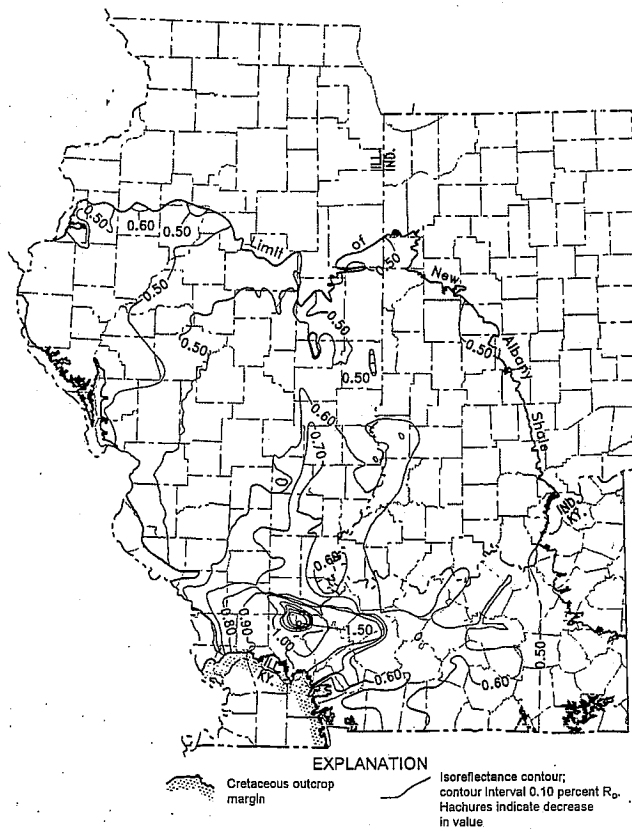


Figure 7. Isoreflectance contour map based on random reflectance measurements of vitrinite in the New Albany Shale (modified from Comer and others, 1994).

99.2 percent was calculated for this range of source-rock percentages and was used to determine the volume of source rock in the Blocher Member for the 0.6- to 0.7-percent R_o interval (table 2). The lower source-rock percentage in the cored Blocher Member of the C-174 well was used to calculate the volume of source rock at vitrinite reflectance intervals greater than 0.9 percent R_o (table 2). Source-rock percentages for the 0.7- to 0.8- and 0.8- to 0.9-percent R_o intervals were estimated by interpolating between the 99.2 percent of the 0.6- to 0.7-percent R_o interval and the 87.2 percent of the 0.9- to 1.0-percent R_o interval. Midpoint reflectance values (i.e., 0.75 and 0.8 percent R_o) for the two intermediate vitrinite-reflectance intervals were used in this linear interpolation to derive the source-rock percentages given in table 2. The resulting total source-rock volume for the Blocher Member is 360 km^3 (table 2), with 56 percent of this total occurring in the 0.6- to 0.7-percent R_o interval.

Estimating the source-rock percentage of the Selmier Member is more complicated because of the large variations in amount of non-source rocks and the limited data. Table 3 shows the source-rock percentage varies from 2 to 93 percent, and the mean value of 55 percent would not accurately represent the entire member. A more representative approach was used by estimating source-rock percentages

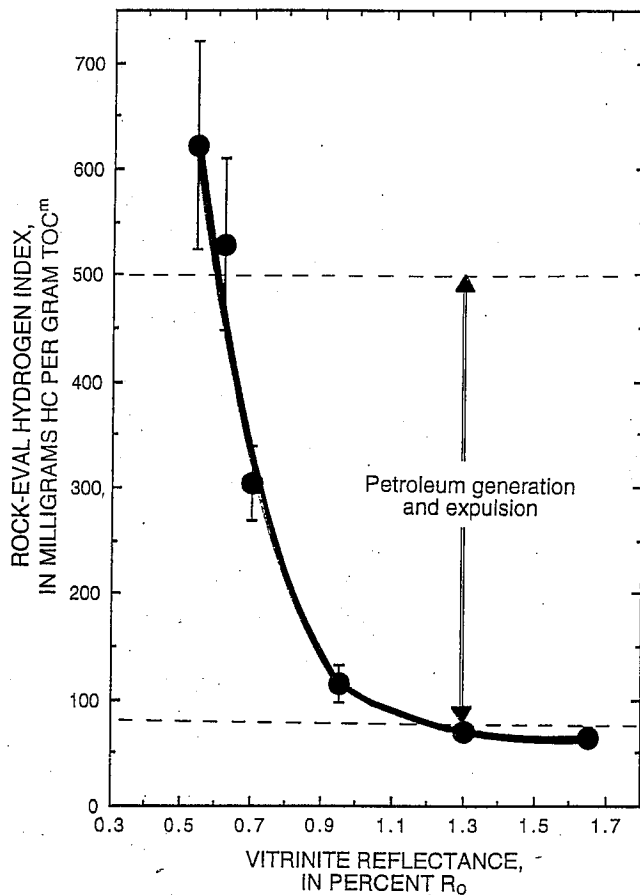


Figure 8. Empirical relationship between mean Rock-Eval hydrogen indices (HI) and mean random vitrinite reflectances (% R_o) from core samples of New Albany Shale at different levels of thermal maturity within the Illinois Basin (see table 1). The relationship from a vitrinite reflectance of 0.5 to 0.95 is described by the expression: $HI = 6,558 \times \exp(-4.27 \times \% R_o)$, and the relationship from a vitrinite reflectance of 0.95 to 1.65 is described by the expression: $HI = 154.69(\% R_o)^2 - 472.92(\% R_o) + 424.36$. HC, hydrocarbons; TOC^m, total organic carbon of mature and over-mature source rocks. Vertical bars indicate one standard deviation.

for the Selmier Member in Illinois and Indiana from cores in the 01IN, 13IL, and 11IL wells and in Kentucky from cores in the 01KY, C-174, and 11IL wells. In Illinois and Indiana, the source-rock percentage of 92.6 percent (table 2) for the 0.6- to 0.7-percent R_o interval is based on the value derived from the 01IN core (table 3). Source-rock percentages for the vitrinite reflectance intervals between 0.7 and 1.3 percent R_o in Illinois and Indiana are derived from a linear interpolation between values from cores in wells 13IL and 11IL. The interpolated values were determined for the midpoint vitrinite reflectance values of each interval (i.e., 0.75, 0.85, 0.95, 1.05, 1.15, and 1.25 percent R_o), and the resulting source-rock percentages of the vitrinite reflectance intervals higher than 1.3 percent R_o are kept constant at 51.8 percent on the basis of the 11IL core.

The source-rock percentages for the Selmier Member in Kentucky are calculated in a similar manner, but with core data from wells 01KY, C-174, and 11IL (table 3). Values in table 3 for midpoints of vitrinite reflectance intervals between 0.6 and 0.9 percent R_o are derived by a linear interpolation between core data of wells 01KY and C-174. The source-rock percentage of 54.6 percent of the C-174 core is used for the 0.9- to 1.0-percent R_o interval. The value of 53.8 percent for the 1.0- to 1.1-percent R_o interval is derived from an interpolation between values from wells C-174 and 11IL. The total volume of source rock in the Selmier Member in all three States is 599 km³ (table 2), with 58 percent of the total occurring in the 0.6- to 0.7-percent R_o interval. The interpolations used to determine the source-rock percentage of the Selmier Member rely on a positive relationship between source-rock thickness and thermal maturity. Although a general relationship of this type is observed by comparing rock thickness in figure 5 with thermal maturity in figure 7, the relationship is not a simple linear relationship as assumed in these interpolations. More work is needed to better constrain these source-rock percentages, but, based on available data, these values represent the best estimates currently available.

ORGANIC RICHNESS

More than 90 volume percent of the organic matter in source rocks of the New Albany Shale consists of oil-prone, amorphous, type-II kerogen (Barrows and Cluff, 1984). The amount of organic matter within the source rocks is most readily expressed as weight percent of total organic carbon (TOC). Cores from six wells were used to determine a weighted mean TOC value for source rocks in each member. These cores represent complete sections and were sampled in sufficient detail to provide representative weighted means, which are given in table 4. The means were weighted with respect to the thickness of source rock represented by the sample analyzed for TOC. The low thermal maturity level (<0.6 percent R_o) of these cores ensures that no significant loss of organic carbon has occurred due to expulsion of generated petroleum or natural gas. The remaining challenge is how to project these weighted TOC means into the higher thermal maturity levels of the basin where organic carbon is reduced due to expulsion of petroleum and natural gas and where source-rock thicknesses may increase by two-fold (table 3).

Two simple models for determining the original TOC content of the thermally mature and over-mature source rocks were used in this study. These two models are considered to represent end members, with the original TOC content most likely occurring somewhere between them. The first model assumes that the amount of organic matter deposited and preserved in source rocks remained constant during the deposition of the source rocks over the entire

Table 1. Summary of TOC and Rock-Eval data on cores of New Albany Shale representing different levels of thermal maturity in the Illinois Basin.

[Mean and standard deviation are given for TOC and Rock-Eval data. R_o , mean random vitrinite reflectance; TOC, total organic carbon; Rx, rock; wt. %, weight percent. See text under heading of Rock-Eval Pyrolysis for definition of S_1 and S_2 hydrocarbons]

Wells included*	Number of samples	R_o (percent)	TOC (wt. % Rx)	S_1 hydrocarbons (mg/g Rx)	S_2 hydrocarbons (mg/g Rx)	Hydrogen index (mg S_2 /TOC)
SDH-305, SDH-308 SDH-309, 02IL	30	0.54	7.07± 3.96	1.75±0.52	42.25±21.33	621.6±98.3
01IN, C-219	10	0.62	4.66±2.30	1.29±0.81	24.05±9.14	528.8±81.7
13IL	5	0.70	4.26±1.24	1.44±0.30	13.10±4.22	304.2±34.7
C-174	9	0.95	4.16±1.46	0.30±0.17	4.76±1.72	114.7±17.6
11IL	2	1.30	5.48±2.00	0.59±0.24	4.00±1.67	71.0±5.0
RM	5	1.65	6.91±1.44	0.34±0.11	4.66±1.65	65.2±8.8

* SDH-305, Indiana Geological Survey No. 305, Jackson Co., Indiana.
 SDH-308, Indiana Geological Survey No. 308, Harrison Co., Indiana.
 SDH-309, Indiana Geological Survey No. 309, Floyd Co., Indiana.
 02IL, Tristar Production No. 1D, Lancaster, Effingham Co., Illinois.
 01IN, Phegley Farms No. 1, Sullivan Co., Indiana.
 C-219, Canton Rock Quarry, Trigg Co., Kentucky.
 13IL, Gordon T. Jenkins No. 1 Simpson, Wayne Co., Illinois.
 C-174, Sheridan Target D.D.H. 66ST-4, Crittenden Co., Kentucky.
 11IL, Rector & Stone No. 1 Missouri Portland Cement, Hardin Co., Illinois.
 RM, Rose Mineral Exploration, Hardin Co., Illinois.

basin, and the thickening of source rocks is the result of increased input of inorganic components. This dilution model predicts that the weighted mean TOC content of a source rock would decrease proportionally to increases in thickness of a source rock. The second model assumes that the amount of organic matter deposited and preserved in a source rock maintains constant proportionality with input of inorganic components. This steady-state model predicts that the weighted mean TOC content of a source rock would remain constant irrespective of changes in the thickness of a source rock.

The dilution model requires determining the fixed mass of organic carbon deposited and preserved beneath a surface-area unit of each source-rock member. Table 4 gives the calculated mass of organic carbon per square meter of source-rock surface area for each source-rock member in the thermally immature wells. The bulk-rock densities (ρ_{Rx}) needed for this determination were calculated from a linear relationship established between TOC values and bulk-rock densities measured on six immature samples of New Albany Shale, with TOC values ranging from 2.46 to 14.34 weight percent and bulk densities ranging from 2.15 to 2.37 Mg/m³. Using the relationship, $\rho_{Rx} = -0.018(\text{TOC}) + 2.40$, mean bulk-rock densities were determined from the weighted mean TOC values of each member represented in the well cores (table 4). Multiplying these mean bulk densities by the thickness of source rock yields the mass of source rock

beneath 1 m² of surface area for the source-rock members in each well. The mass of organic carbon deposited and preserved beneath 1 m² of surface area may then be determined by multiplying the source-rock mass per square meter by the decimal fraction of the weighted mean TOC content. Large standard deviations (table 4) indicate that the assumption of this model is questionable and the use of mean values is difficult to justify. An alternative is to use TOC mass/area units from a thermally immature well that is nearest to the thermally maturing depocenter. The well used for this purpose was 01KY, which is less than 1 km from an isorelectance contour of 0.6 percent R_o and 72 km from an isorelectance contour of 1.0 percent R_o . The calculated values of TOC mass/area for this well are 0.82, 2.38, and 4.17 Mg/m² for the Blocher, Selmier, and upper shale composite members, respectively (table 4). Multiplying these values by the surface areas given for each member at 0.1-percent R_o intervals (table 2) gives, according to this model, the original TOC mass of thermally mature and over-mature source rocks (table 5).

The steady-state model requires determining the mean mass of organic carbon deposited and preserved within a volume unit of source rock. Figure 9 shows the relationship between mass of TOC/surface area and thickness of source rock. A linear regression of these data from tables 3 and 4 gives a slope of 197.7 kg of TOC per cubic meter of source rock and a near-zero intercept for all source-rock members.

Table 2. Surface area and volume of source rocks in members of of New Albany Shale at 0.1-percent R_o intervals.

R_o interval (percent)	Surface area (km^2)	Total volume (km^3)	Source-rock percentage	Source-rock volume (km^3)
Upper shale composite.				
0.6-0.7	24,603.20	695.86	97.5 ¹	678.46
0.7-0.8	5,698.14	169.75	97.5 ¹	165.51
0.8-0.9	2,350.86	56.71	97.5 ¹	55.29
0.9-1.0	4,193.75	149.69	97.5 ¹	145.95
1.0-1.1	1,294.78	52.73	97.5 ¹	51.41
1.1-1.2	179.23	6.89	97.5 ¹	6.71
1.2-1.3	149.13	5.86	97.5 ¹	5.72
1.3-1.4	107.75	4.40	97.5 ¹	4.29
1.4-1.5	139.96	5.61	97.5 ¹	5.47
>1.5	12.07	0.49	97.5 ¹	0.48
				Total = 1,119.29
Selmier Member				
Indiana/Illinois:				
0.6-0.7	14,476.81	201.08	92.6 ²	186.20
0.7-0.8	4,915.35	91.24	61.2 ³	56.63
0.8-0.9	1,484.26	17.02	59.5 ³	10.13
0.9-1.0	2,158.21	54.91	57.8 ³	31.74
1.0-1.1	814.55	38.18	56.1 ³	21.42
1.1-1.2	179.24	8.25	54.4 ³	4.49
1.2-1.3	149.13	6.85	52.7 ³	3.61
1.3-1.4	107.75	5.39	51.8 ⁴	2.79
1.4-1.5	139.95	7.05	51.8 ⁴	3.65
>1.5	12.07	0.66	51.8 ⁴	0.34
Kentucky:				
0.6-0.7	10,126.43	209.69	76.6 ⁵	160.62
0.7-0.8	782.77	31.09	69.3 ⁵	21.55
0.8-0.9	866.59	37.92	61.9 ⁵	23.47
0.9-1.0	2,035.43	105.81	54.6 ⁶	57.77
1.0-1.1	480.24	26.25	53.8 ⁷	14.12
				Total = 598.57
Blocher Member				
0.6-0.7	24,603.22	203.70	99.2 ⁸	202.07
0.7-0.8	5,698.15	47.76	96.8 ⁹	46.23
0.8-0.9	2,350.85	22.12	92.0 ⁹	20.35
0.9-1.0	4,193.76	65.48	87.2 ¹⁰	57.09
1.0-1.1	1,294.78	28.07	87.2 ¹⁰	24.47
1.1-1.2	179.23	3.64	87.2 ¹⁰	3.17
1.2-1.3	149.13	2.24	87.2 ¹⁰	1.95
1.3-1.4	107.75	2.44	87.2 ¹⁰	2.13
1.4-1.5	139.96	3.01	87.2 ¹⁰	2.62
>1.5	12.07	0.25	87.2 ¹⁰	0.22
				Total = 360.31

¹Mean value of upper shale composite given in table 3 for all wells.²Values from 01IN well in table 3.³Value from linear interpolation between 13IL and 11IL wells in table 3.⁴Value from 11IL well in table 3.⁵Value from linear interpolation between 01KY and C-174 wells in table 3.⁶Value from C-174 well in table 3.⁷Value from linear interpolation between C-174 and 11IL wells in table 3.⁸Mean value of Blocher Member (excluding C-174 well) in table 3.⁹Value from linear interpolation between the mean value of 99.2 percent in the 0.6- to 0.7-percent R_o interval and 87.2 percent in the C-174 well of table 3.¹⁰Value from C-174 well in table 3.

Table 3. Thickness and percentage of source rock in members of New Albany Shale determined in available well cores.[Mean value for source-rock percentage for Blocher Member does not include data for well C-174; R_o , mean random vitrinite reflectance]

Well name	R_o (percent)	Total thickness (m)	Source-rock thickness (m)	Source-rock percentage	Source of data
Upper shale composite					
SDH-308	0.44	22.11	21.75	98.4	1
SDH-273	0.46	18.03	17.96	99.6	1
A-11	0.50	16.29	15.68	96.3	2
SDH-305	0.50	28.46	27.24	95.7	1
SDH-309	0.53	26.29	26.05	99.1	1
01KY	0.56	22.17	20.47	92.3	3
01IN	0.63	25.96	25.22	97.1	1
13IL	0.70	40.66	40.66	100.0	3
C-174	0.95	36.11	35.81	99.2	4
				Mean = 97.5	
Selmier Member					
SDH-273	0.46	8.69	3.48	40.0	1
SDH-305	0.50	4.70	0.09	1.9	1
01KY	0.56	16.41	13.66	83.2	3
01IN	0.63	9.03	8.36	92.6	1
13IL	0.70	18.85	11.71	62.1	3
C-174	0.95	62.53	34.16	54.6	4
11IL	1.30	53.50	27.72	51.8	3
				Mean = 55.2	
Blocher Member					
SDH-308	0.44	4.82	4.82	100.0	1
SDH-273	0.46	4.97	4.94	99.4	1
A-11	0.50	3.48	3.48	100.0	2
SDH-305	0.50	3.66	3.60	98.4	1
SDH-309	0.53	5.12	5.06	98.8	1
01KY	0.56	5.82	5.75	98.8	3
01IN	0.63	3.23	3.17	98.1	1
13IL	0.70	8.78	8.78	100.0	3
				Mean = 99.2	
C-174	0.95	14.03	12.23	87.2	4

¹Nancy Hasenmueller, Indiana Geological Survey.²Thomas Robl, Center for Applied Energy Research, University of Kentucky.³Cluff and others (1981) and Frost (1980).⁴John W. Smith, Laramie Petroleum Research Center.

This slope may be equated to a mean TOC content of 8.82 weight percent with a bulk-rock density of 2.24 Mg/m³ for all the source rocks in the New Albany Shale. Multiplying this slope value by the source-rock volumes given for each member at 0.1-percent R_o intervals (table 2) gives, according to this model, the original TOC mass in the mature and over-mature source rocks (table 5). Comparison of results from both models (table 5) shows the TOC mass determined by the steady-state model is 44 percent greater than that determined by the dilution model.

A model predicting an increase in TOC content of the source rocks as the New Albany Shale thickens into the depocenter was not considered likely on the basis of available

data and analogous studies of source rocks in other basins. A significant portion of this thickening is a result of an increase in thickness of greenish or olive-gray, organic-lean shales as described in cores from wells 11IL and 13IL (Cluff and others, 1981). This increase in non-source rocks is especially obvious in the Selmier Member, as suggested by the thicknesses of indistinctly bedded and poorly laminated shales in the cross sections of figure 3. These rocks are typically partially bioturbated and are considered not to represent periods of anoxic conditions within the basin (Cluff, 1980). In addition, thickening of Devonian-Mississippian source rocks around the Findlay arch and in the western Appalachian Basin results in a decrease in their TOC contents. Robl and

Table 4. Organic richness data from complete cores of thermally immature source-rock members within the New Albany Shale.

[Sources of data are same as those footnoted in table 3. TOC, total organic carbon; wt.%, weight percent; SRx, source rock]

Well name ¹	Weighted mean TOC ² (wt. % SRx)	TOC range (wt. % SRx)	Number of analyses	Bulk-rock density ³ (Mg/m ³)	SRx mass/surface area ⁴ (Mg/m ²)	TOC mass/surface area ⁵ (Mg/m ²)
Upper shale composite						
SDH-308	7.92	2.98-15.00	25	2.26	49.155	3.893
SDH-273	8.62	5.19-11.30	13	2.24	40.230	3.468
A-11	10.12	5.52-15.95	27	2.22	34.810	3.523
SDH-305	9.11	3.80-16.10	21	2.24	61.018	5.559
SDH-309	8.07	2.68-16.20	32	2.25	58.613	4.730
01KY	9.10	5.81-14.03	4	2.24	45.853	4.173
					Mean	4.224
					Standard deviation	±0.732
Selmier Member						
SDH-273	8.14	7.85-8.28	3	2.25	7.830	0.637
01KY	7.71	3.40-11.95	8	2.26	30.872	2.380
					Mean	1.509
					Standard deviation	±0.871
Blocher Member						
SDH-308	9.37	5.73-11.40	6	2.23	10.749	1.007
SDH-273	5.96	2.51-11.76	7	2.29	11.313	0.674
A-11	9.07	6.03-12.19	5	2.24	7.795	0.707
SDH-305	8.21	7.32-9.70	4	2.25	8.100	0.665
SDH-309	7.81	3.69-9.26	6	2.26	11.436	0.893
01KY	6.25	5.67-7.71	4	2.29	13.168	0.823
					Mean	0.795
					Standard deviation	±0.126

¹Notations for wells SDH-305, SDH-308, and SDH-309 are the same as in the footnote for table 1. 01KY, Orbit Gas No. 1 Ray Clark, Christian Co., Kentucky; SDH-273, Indiana Geological Survey No. 273, Marion Co., Indiana; A-11, Nelson Co., Kentucky.

²Weighted means are calculated by multiplying the TOC value by the thickness of source rock it represents, summing these products for each member, and dividing the sum by the total thickness of source rock for each member.

³Bulk-rock densities (ρ_{Rx}) calculated from a linear expression (i.e., $\rho_{Rx} = -0.0189(\text{TOC}) + 2.40$) based on six samples of New Albany Shale with TOC values ranging from 2.46 to 14.34 weight percent and bulk-rock densities ranging from 2.15 to 2.37 Mg/m³.

⁴Source-rock mass per square meter of surface area is determined by multiplying the bulk-rock density by the source-rock thickness given in table 3.

⁵TOC mass of source rocks per square meter of surface area is determined by multiplying the source-rock mass per square meter of surface area by the decimal fraction of the weighted mean TOC value.

others (1983) report decreases in TOC contents for the Cleveland Member of the Ohio Shale around the Findlay arch from 15 to 9 weight percent with increases in thicknesses from 4 to 18 m, respectively. Similarly, Schmoker (1980) reports decreases in TOC contents for Devonian organic-rich facies in the western Appalachian Basin from 16 to 5 volume percent with increases in thicknesses from 65 to 240 m. A similar inverse relationship between the TOC content and thickness of source rock was reported by Comer (1991, 1992) for Upper Devonian black shales in the Delaware and Anadarko Basins and by Nixon (1973) for the Mowry Shale (Lower Cretaceous) in the western Early Cretaceous Boreal seaway.

EXPULSION FACTOR

The expulsion factor is defined as the amount of petroleum that is expelled from a petroleum source rock as it thermally matures within a sedimentary basin. The unit of this parameter is in volume or mass of petroleum per unit mass of total organic carbon in a source rock. It is important that this parameter not be confused with hydrocarbon generation potential, which is a measure of hydrocarbons generated within rather than expelled from a source rock. Expulsion factors may be determined through evaluation of natural maturation data or laboratory pyrolysis studies. Although useful qualitative concepts and some constraints are obtained from

Table 5. Mass of total organic carbon in New Albany Shale source rocks as determined by dilution and steady-state models.

[TOC, total organic carbon. SS, TOC mass for steady-state model; D, TOC mass for dilution model. Mbr, Member; Fm, Formation]

R _o interval (percent)	Dilution model TOC mass ¹ (Tg)	Steady-state model TOC mass ² (Tg)	Percent difference $\{[(SS/D) \times 100] - 100\}$
Upper shale composite			
0.6-0.7	102,595	134,132	31
0.7-0.8	23,761	32,721	38
0.8-0.9	9,803	10,931	12
0.9-1.0	17,488	28,854	65
1.0-1.1	5,399	10,164	88
1.1-1.2	747	1,327	78
1.2-1.3	622	1,131	82
1.3-1.4	449	848	89
1.4-1.5	584	1,081	85
>1.5	50	95	90
"Mbr" total	161,498	221,284	37
Selmier Member			
0.6-0.7	58,556	68,566	17
0.7-0.8	13,562	15,456	14
0.8-0.9	5,595	6,643	19
0.9-1.0	9,981	17,696	77
1.0-1.1	3,082	7,026	128
1.1-1.2	427	888	108
1.2-1.3	355	714	101
1.3-1.4	256	552	116
1.4-1.5	333	722	117
>1.5	29	67	131
Mbr total	92,176	118,329	28
Blocher Member			
0.6-0.7	20,175	39,950	98
0.7-0.8	4,672	9,140	96
0.8-0.9	1,928	4,023	109
0.9-1.0	3,439	11,288	228
1.0-1.1	1,062	4,837	355
1.1-1.2	147	627	327
1.2-1.3	122	385	216
1.3-1.4	88	420	377
1.4-1.5	115	518	350
>1.5	10	44	340
Mbr total	31,758	71,232	124
Fm total	285,432	410,845	44

¹Surface areas given in table 2 for the upper shale composite, Selmier, and Blocher Members are multiplied by 4.17, 2.38, and 0.82 Mg/m², respectively.

²Source-rock volumes given in table 2 are multiplied by the 197.7 × 10⁶ Mg/km³ slope derived from linear regression of data plotted in figure 9.

expulsion factors derived from natural maturation data (Momper, 1978; Jones, 1981), the vastness of sedimentary basins, limited subsurface data, and dynamics of fluid flow

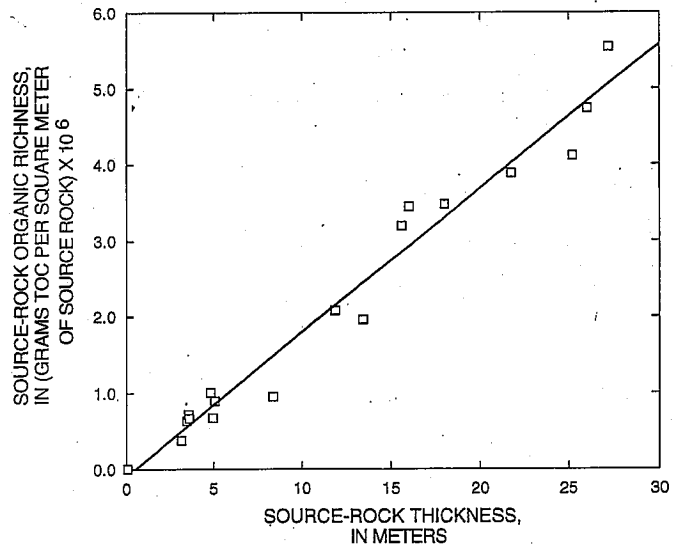


Figure 9. Relationship between grams of total organic carbon (TOC) per square meter of source-rock surface area (table 4) and source-rock thickness, in meters (table 3), for thermally immature source rocks. Linear regression describes the relationship with a zero intercept and a slope of 197.7 kg TOC/m³ (correlation coefficient = 0.991).

over geological time make it unfeasible to obtain applicable quantitative concepts. Therefore, it becomes essential to employ laboratory pyrolysis methods that simulate the natural processes to develop a quantitative determination of expulsion factors. The two laboratory pyrolysis methods considered in this study are Rock-Eval pyrolysis and hydrous pyrolysis.

ROCK-EVAL PYROLYSIS

This open anhydrous-pyrolysis method typically consists of heating 50 to 150 mg of powdered rock in an oven that is continuously swept with flowing helium at low pressures (<138 kPa). The powdered rock is initially heated at a constant temperature between 250°C and 300°C for 3 to 5 minutes. The volatilized products that are released during this heating are swept with helium into a flame ionization detector for quantification of hydrocarbon content. These readily volatilized hydrocarbons are considered to be unbound to the organic matter within the rock and are designated as "S₁" hydrocarbons. Immediately following this initial isothermal heating, the powdered rock is heated at a rate of 25°C/min or 30°C/min to a final temperature of 600°C, which is maintained for 1 minute. The pyrolyzed products that are generated during this heating are vaporized and swept with helium into a flame ionization detector for quantification of hydrocarbon content. These pyrolyzed hydrocarbons are considered to be bound to the organic matter within the rock and are designated as "S₂" hydrocarbons.

Although vaporization of generated hydrocarbons by this pyrolysis method has no resemblance to natural conditions under which petroleum is generated and expelled in sedimentary basins (Lewan, 1993, 1994), it has been successfully used to provide rapid evaluations of the relative hydrocarbon potential and thermal maturity of source rocks (Bordenave and others, 1993). Cooles and others (1986) extended these routine screening applications of Rock-Eval pyrolysis to determine expulsion factors of source rocks for direct application in petroleum material-balance studies within sedimentary basins (e.g., Mackenzie and Quigley, 1988; MacGregor and McKenzie, 1987). The central assumption in this algebraic scheme is that a fixed fraction of organic carbon cannot be vaporized as S_1 and S_2 hydrocarbons by Rock-Eval pyrolysis irrespective of the thermal maturity of a source rock. This inactive organic carbon is referred to as being "inert." Another assumption in this approach is that gas generation is sufficiently negligible during petroleum formation to allow hydrocarbon yields to be equated to petroleum yields (Cooles and others, 1986). This approach requires determination of total, pyrolyzed, and volatilized organic carbon in immature and mature source rocks. The volatilized and pyrolyzed organic carbon (C_1 and C_2 , respectively) are determined respectively by multiplying the S_1 and S_2 hydrocarbons from Rock-Eval pyrolysis by 0.85, which is an approximation of the carbon fraction of petroleum (Cooles and others, 1986). Inert organic carbon (C_I) is calculated by subtracting the sum of C_1 and C_2 from the total organic carbon (TOC) of a source rock. An "o" superscript is used to denote organic carbon from immature source rocks that have not lost organic carbon due to petroleum or natural gas expulsion. An "m" superscript is used to denote organic carbon from mature or over-mature source rocks that have lost organic carbon due to petroleum or natural gas expulsion. All of these organic carbon values are in units of grams of carbon per gram of source rock (g C/g Rx).

The first step in this approach is to determine the fraction of total organic carbon that occurs as inert organic carbon in the immature source rock (i.e., X_I^o), which is determined by the expression:

$$X_I^o = \frac{TOC^o - (C_2^o + C_1^o)}{TOC^o} \quad (1)$$

The assumption that this fraction of inert organic carbon remains constant with increasing thermal maturity (i.e., $X_I^o = X_I^m$) allows the calculation of the original total organic carbon of the immature equivalent of a mature source rock (TOC^{om}) by the expression:

$$TOC^{om} = \frac{TOC^m - (C_2^m + C_1^m)}{X_I^o} \quad (2)$$

The difference between the total organic carbon calculated for the original source rock (TOC^{om}) and measured on the mature source rock (TOC^m) is considered the amount of organic carbon lost from the mature source rock as a result of hydrocarbon expulsion. This quantity may be divided by the original total organic carbon of the mature source rock (TOC^{om}) to determine the grams of expelled carbon per gram of original total organic carbon of the source rock. Dividing this quantity by the carbon fraction of petroleum (0.85) gives an expulsion factor in grams of petroleum expelled from a mature source rock per gram of original total organic carbon in its original immature state (g HC/g TOC^{om}).

This approach is applied to the mean values of the natural data set given in table 1. The fraction of inert organic carbon (X_I^o) is calculated by equation 1 to be 0.4712. This value is used to calculate the original total organic carbon of the mature and over-mature source rocks using equation 2. The resulting Rock-Eval expulsion factors are given in table 6. The relation between these expulsion factors and hydrogen indices may be described by a second-degree polynomial as shown in figure 10. This expression and the relation between percent R_o values and hydrogen indices in figure 9 allow the calculation of expulsion factors for the midpoint values for each 0.1-percent R_o interval. Results of this operation and the calculated quantities of expelled hydrocarbons for the dilution and steady-state models are given in table 7. Assuming that the expelled hydrocarbons measured by Rock-Eval pyrolysis are similar to petroleum produced in the Illinois Basin, barrels of expelled petroleum are calculated on the basis of an API gravity of 36.8° (Mast, 1970). The dilution and steady-state models respectively predict 786 and 1,172 billion barrels of petroleum were expelled from source rocks in the New Albany Shale.

HYDROUS PYROLYSIS

This closed pyrolysis method typically consists of heating 300 to 500 g of rock chips (0.5 to 2.0 cm) in the presence of liquid water in a 1-L reactor at a constant temperature between 250°C and 365°C for 3 to 6 days (Lewan, 1993). Internal pressures are a result of water vapor and generated gases within the headspace of the reactor and may range from 10 to more than 20 MPa, depending on the experimental temperature, organic richness of sample, and heating duration. If the rock sample is a potential source rock and the experimental time and temperature conditions are sufficient, a free-flowing liquid oil is expelled from the rock and accumulates on the surface of the water within the reactor (Lewan, 1985; Huizinga and others, 1988; Peters and others, 1990; Baskin and Peters, 1992). This expelled oil has a gross, molecular, and isotopic composition similar to natural petroleum (Lewan, 1983, 1993; Peters and others, 1990; Lewan and others, 1978, 1983). API gravities of these oils

Table 6. Rock-Eval expulsion factors calculated from geochemical data given in table 1 using equations 1 and 2.

[R_o , mean random vitrinite reflectance; C, carbon; Rx, rock; HC, hydrocarbons; TOC^m , total organic carbon in mature or over-mature source rock; C_1^m , volatilized organic carbon determined by Rock-Eval pyrolysis of mature or over-mature source rocks ($S_1 \times 0.85$); C_2^m , pyrolyzed organic carbon determined by Rock-Eval pyrolysis of mature or over-mature source rocks ($S_2 \times 0.85$); C_I^m , inert carbon in mature or over-mature rock ($TOC^m - [C_1^m + C_2^m]$); TOC^{om} , original immature total organic carbon of mature or over-mature source rock (see eq. 1); ($TOC^{om} - TOC^m$) is equal to the amount of organic carbon expelled from mature or over-mature source rock; expelled carbon is equal to $(TOC^{om} - TOC^m)/TOC^{om}$; expelled hydrocarbons are equal to expelled carbon divided by 0.85]

	Wells included (see footnote)				
	01IN C-219	13IL	C-174	11IL	RM
R_o (percent)	0.62	0.70	0.95	1.30	1.65
TOC^m (g C/g Rx)	0.0466	0.0426	0.0416	0.0548	0.0691
C_1^m (g C/g Rx)	0.0011	0.0012	0.0003	0.0005	0.0003
C_2^m (g C/g Rx)	0.0204	0.0111	0.0040	0.0034	0.0040
C_I^m (g C/g Rx)	0.0251	0.0303	0.0373	0.0509	0.0648
TOC^{om} (g C/g Rx)	0.0533	0.0643	0.0792	0.1080	0.1375
$TOC^{om} - TOC^m$ (g C/g Rx)	0.0067	0.0217	0.0376	0.0532	0.0684
Expelled carbon (g C/g TOC^{om})	0.1257	0.3375	0.4747	0.4926	0.4975
Expelled HC (g HC/g TOC^{om})	0.148	0.397	0.558	0.580	0.585

Well notation given in footnote for table 1.

may vary from 25° to 42°, depending on the type of organic matter within the rock and the experimental conditions. These close similarities between hydrous-pyrolysis oils and natural petroleum suggest that, despite the higher temperatures used in hydrous pyrolysis, the mechanisms responsible for natural petroleum generation and expulsion are being simulated in the experiments.

A series of hydrous-pyrolysis experiments were conducted on aliquots of a thermally immature sample of New Albany Shale to determine expulsion factors at different maturity levels. The sample (931026-3) was collected from the Clegg Creek Member at a roadcut along Highway 311, west of Sellersburg, Ind. This sample is organic rich and thermally immature, with a vitrinite reflectance less than 0.5 percent R_o , a T_{max} of 425°C, a production index of 0.03, a hydrogen index of 604 mg/g TOC, and a TOC of 14.3 weight percent. Care was taken to collect a blocky unweathered sample, which was crushed to gravel size (0.5 to 2.0 cm) for the experiments. Each of the experiments consisted of heating 300 g of crushed rock and 400 g of distilled water in 1-L reactors for 72 hours at constant temperatures between 270°C and 365°C. The expelled oil was collected from the reactor as described by Lewan (1993). The quantities of expelled oil are calculated as mass of expelled oil per mass of original TOC (i.e., TOC^o). These expulsion factors increase with increasing experimental temperature, with the exception of one outlying data point (fig. 11).

Relating these hydrous-pyrolysis expulsion factors to the 0.1-percent R_o intervals of the source rocks is

accomplished through the relationships both of these parameters have with the hydrogen index parameter determined by Rock-Eval pyrolysis of natural and experimentally matured samples. First, the relationship established between percent R_o and hydrogen index for naturally matured samples (fig. 8) is used to determine the hydrogen index for the midpoint percent R_o values for each 0.1-percent R_o interval. Secondly, these calculated hydrogen indices are used in the relationship between hydrogen indices and expulsion factors for experimentally matured samples (fig. 12) to determine the hydrous-pyrolysis expulsion factor for each 0.1-percent R_o interval. The results are shown in table 8 along with the quantities of oil expelled for mature and over-mature source rocks for the dilution and steady-state models. Similar to the volume quantities determined from the Rock-Eval expulsion factors in table 7, barrels of oil expelled are calculated on the basis of a 36.8° API-gravity oil. The total quantities of expelled petroleum as determined by hydrous pyrolysis are 305 to 473 billion barrels, which represent only 39 to 40 percent of the 786 to 1,172 billion barrels determined by Rock-Eval expulsion factors (table 7).

ACCOUNTABLE IN-PLACE PETROLEUM

Mast and Howard (1990) estimated, on the basis of production data and a recovery factor of 35 percent that the original in-place petroleum for known fields in the Illinois Basin is 12 billion barrels. Most of this in-place petroleum

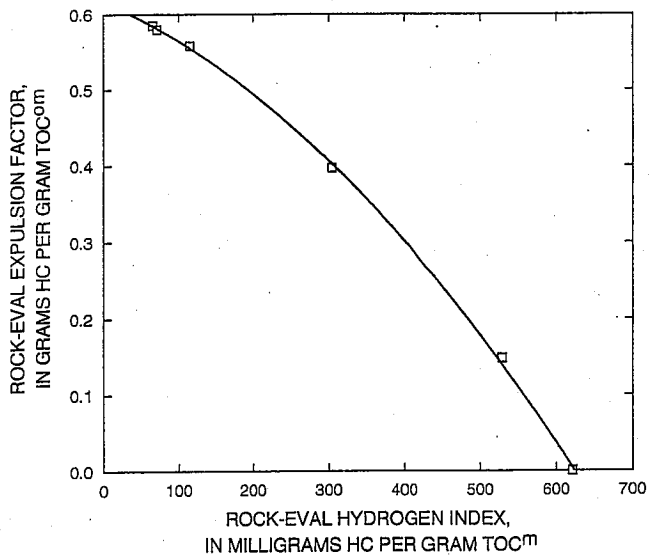


Figure 10. Relationship between expulsion factors (EF) determined by Rock-Eval pyrolysis (table 6) and Rock-Eval hydrogen indices (HI) (table 1). The curve is defined by the expression: $EF = 0.616 - 43.26 \times 10^{-5}(HI) - 88.68 \times 10^{-8}(HI)^2$. HC, hydrocarbons; TOC^{om} , original immature total organic carbon of mature and over-mature source rocks; TOC^m , total organic carbon of of mature and over-mature source rocks.

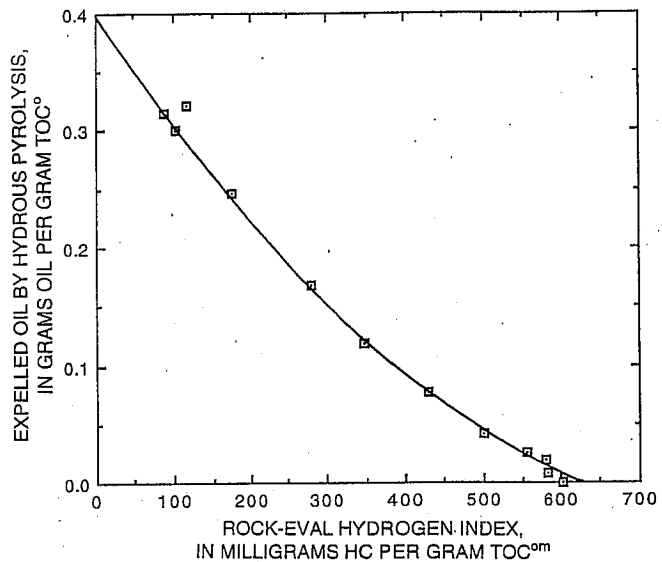


Figure 12. Relationship between expulsion factors derived from hydrous pyrolysis (HPEE) and hydrogen index (HI) determined by Rock-Eval pyrolysis on rock recovered from the hydrous-pyrolysis experiments (fig. 11). The curve is described by the expression: $HPEE = 0.399 - 99.60 \times 10^{-5}(HI) + 573.26 \times 10^{-9}(HI)^2$. HC, hydrocarbons; TOC^{om} , original immature total organic carbon of mature and over-mature source rocks; TOC^o , total organic carbon of immature source rocks.

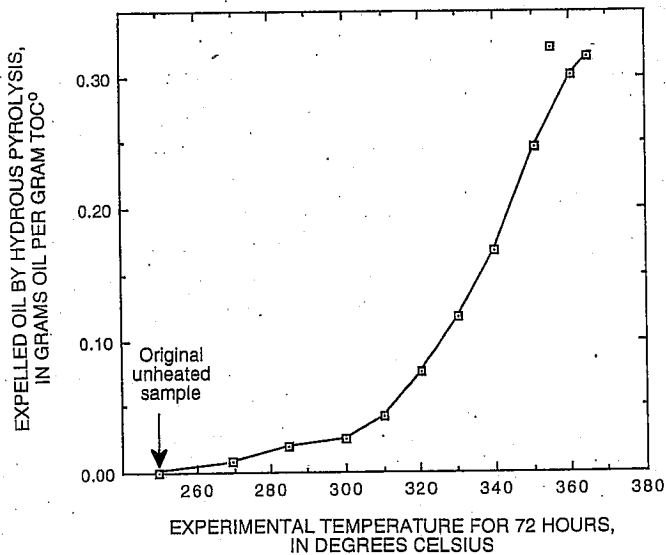


Figure 11. Quantity of petroleum expelled from hydrous-pyrolysis experiments conducted isothermally at temperatures from 270°C to 365°C for 72 hours with aliquots of a sample from Clegg Creek Member of the New Albany Shale. TOC^o , total organic carbon of immature source rocks.

occurs in reservoirs of Carboniferous age, with more than 60 percent occurring in Chesterian sandstones, 18 percent occurring in Valmeyeran sandstones, and 13 percent occurring in Pennsylvanian sandstones. Organic geochemical correlations indicate that more than 99 percent of the

discovered in-place petroleum of the basin has been sourced by the New Albany Shale (Hatch and others, 1990).

In addition to free-flowing producible petroleum, tar-sand deposits occur at or near the surface in various sandstones of Late Mississippian and Early Pennsylvanian age. These tar deposits are dispersed around the margins of the eastern Moorman syncline of western Kentucky and appear to be closely associated with faults, paleovalleys, and stratigraphic traps (Noger, 1987). Conventional oil production has been obtained locally from the same sandstones a few kilometers down-dip (Williams and others, 1982). Some of these deposits have local seeps that flow as much as 10 barrels per year (McGrain, 1976). Resource assessments of these deposits are classified as measured and speculative. Measured reserves are based on core analyses of tar encountered in wells, and speculative reserves are based on reported tar shows on drilling logs, occasional core analyses, and geological interpretations (Noger, 1987). Measured in-place tar has been determined to be 1.73 billion barrels, and speculative in-place tar has been determined to be 1.68 billion barrels (Kuuskraa, 1984; Crysdale and Schenk, 1988; Noger, 1987). These values give a total in-place tar of 3.41 billion barrels, which extends over 2,500 km² of surface area and from the surface to a 183-m depth (Noger, 1987). No organic geochemical correlations have established the source rock for these tars, and opinions vary: Russell (1933) suggests that Late Mississippian marine shales are the source, and

Table 7. Quantity of petroleum expelled from the New Albany Shale as determined by Rock-Eval expulsion factors.[R_o, mean random vitrinite reflectance; HC, hydrocarbon; TOC, total organic carbon; bbl, barrel; Mbr, Member; Fm, Formation]

R _o interval (percent)	Rock-Eval expulsion factor ¹ (kg HC/Mg TOC)	Dilution model		Steady-state model	
		Vaporized hydrocarbons ² (Tg)	Expelled petroleum ⁴ (billion bbl)	Vaporized hydrocarbons ³ (Tg)	Expelled petroleum ⁴ (billion bbl)
Upper shale composite					
0.6-0.7	290.62	29,816	223.07	38,981	291.63
0.7-0.8	437.44	10,394	77.76	14,313	107.08
0.8-0.9	513.91	5,038	37.69	5,617	42.03
0.9-1.0	555.04	9,706	72.62	16,015	119.82
1.0-1.1	565.23	3,052	22.83	5,745	42.98
1.1-1.2	573.09	428	3.20	760	5.69
1.2-1.3	578.95	360	2.69	655	4.90
1.3-1.4	582.86	262	1.96	494	3.70
1.4-1.5	585.06	341	2.55	633	4.73
>1.5	585.55	29	0.22	56	0.42
	"Mbr" total	59,426	444.59	83,269	622.98
Selmier Member					
0.6-0.7	290.62	17,017	127.31	19,926	149.08
0.7-0.8	437.44	5,932	44.38	6,761	50.58
0.8-0.9	513.91	2,875	21.51	3,414	25.54
0.9-1.0	555.04	5,540	41.45	9,822	73.48
1.0-1.1	565.23	1,742	13.03	3,971	29.71
1.1-1.2	573.09	244	1.83	509	3.81
1.2-1.3	578.95	205	1.54	413	3.09
1.3-1.4	582.86	149	1.12	321	2.41
1.4-1.5	585.06	195	1.46	422	3.16
>1.5	585.55	17	0.13	39	0.29
	Mbr total	33,916	253.76	45,598	341.15
Blocher Member					
0.6-0.7	290.62	5,863	43.86	11,610	86.86
0.7-0.8	437.44	2,044	15.29	3,998	29.91
0.8-0.9	513.91	991	7.41	2,068	15.47
0.9-1.0	555.04	1,909	14.28	6,265	46.87
1.0-1.1	565.23	600	4.49	2,734	20.46
1.1-1.2	573.09	84	0.63	359	2.69
1.2-1.3	578.95	71	0.53	223	1.67
1.3-1.4	582.86	52	0.39	245	1.83
1.4-1.5	585.06	67	0.50	303	2.27
>1.5	585.55	6	0.04	26	0.19
	Mbr total	11,687	87.42	27,831	208.22
	Fm total	105,029	785.77	156,698	1,172.34

¹Hydrogen index for the midpoint of each 0.1-percent R_o interval was determined from the relationship in figure 8 and the resulting hydrogen index was used to determine the Rock-Eval expulsion factor from the relationship in figure 10.

²Rock-Eval expulsion factor ($\times 10^{-3}$) is multiplied by the TOC mass as determined by the dilution model in table 5.

³Rock-Eval expulsion factor ($\times 10^{-3}$) is multiplied by the TOC mass as determined by the steady-state model in table 5.

⁴A petroleum density of 840.76 kg/m³ (i.e., 36.8°API gravity) is divided into the mass of vaporized hydrocarbons and the quotient is multiplied by 6.29 to convert cubic meters to billion barrels.

Table 8. Quantity of petroleum expelled from the New Albany Shale as determined by hydrous-pyrolysis expulsion factors.[R_o, mean random vitrinite reflectance; TOC, total organic carbon; bbl, barrel; Mbr, Member; Fm, Formation]

R _o interval (percent)	Hydrous-pyrolysis expulsion factor ¹ (kg oil/Mg TOC)	Dilution model		Steady-state model	
		Expelled oil ² (Tg)	Expelled petroleum ⁴ (billion bbl)	Expelled oil ³ (Tg)	Expelled petroleum ⁴ (billion bbl)
Upper shale composite					
0.6-0.7	87.45	8,972	67.12	11,730	87.75
0.7-0.8	173.94	4,133	30.92	5,691	42.58
0.8-0.9	242.86	2,381	17.81	2,655	19.86
0.9-1.0	292.43	5,114	38.26	8,438	63.13
1.0-1.1	306.76	1,656	12.39	3,118	23.33
1.1-1.2	318.52	238	1.78	423	3.16
1.2-1.3	327.75	204	1.53	371	2.77
1.3-1.4	334.14	150	1.12	283	2.12
1.4-1.5	337.83	197	1.48	365	2.73
>1.5	338.66	17	0.13	32	0.24
	"Mbr" total	23,062	172.53	33,106	247.67
Selmier Member					
0.6-0.7	87.45	5,121	38.31	5,996	44.86
0.7-0.8	173.94	2,359	17.65	2,688	20.11
0.8-0.9	242.86	1,359	10.17	1,613	12.07
0.9-1.0	292.43	2,919	21.84	5,175	38.71
1.0-1.1	306.76	945	7.07	2,155	16.12
1.1-1.2	318.52	136	1.02	283	2.12
1.2-1.3	327.75	116	0.87	234	1.75
1.3-1.4	334.14	86	0.64	184	1.38
1.4-1.5	337.83	112	0.84	244	1.82
>1.5	338.66	10	0.07	23	0.17
	Mbr total	13,163	98.48	18,595	139.11
Blocher Member					
0.6-0.7	87.45	1,764	13.20	3,494	26.14
0.7-0.8	173.94	813	6.08	1,590	11.89
0.8-0.9	242.86	468	3.50	977	7.31
0.9-1.0	292.43	1,006	7.52	3,301	24.70
1.0-1.1	306.76	326	2.44	1,484	11.10
1.1-1.2	318.52	47	0.35	200	1.49
1.2-1.3	327.75	40	0.30	126	0.94
1.3-1.4	334.14	29	0.22	140	1.05
1.4-1.5	337.83	39	0.29	175	1.31
>1.5	338.66	3	0.03	15	0.11
	Mbr total	4,535	33.93	11,502	86.04
	Fm total	40,760	304.94	63,203	472.82

56.6%

32.3%

11.126%

¹Hydrogen indices for the midpoint value of each 0.1-percent R_o interval was determined from the relationship in figure 8 and the resulting hydrogen indices were used to determine hydrous-pyrolysis expulsion factors from the relationship in figure 12.

²Hydrous-pyrolysis expulsion factor ($\times 10^{-3}$) is multiplied by the TOC mass as determined by the dilution model in table 5.

³Hydrous-pyrolysis expulsion factor ($\times 10^{-3}$) is multiplied by the TOC mass as determined by the steady-state model in table 5.

⁴A petroleum density of 840.76 kg/m³ (i.e., 36.8°API gravity) is divided into the mass of expelled oil and the quotient is multiplied by 6.29 to convert cubic meters to billion barrels.

Walters (1974) suggests that Devonian-Mississippian black shales (i.e., New Albany Shale) are the source. With no other proven source rocks younger than the New Albany Shale (Cluff and Byrnes, 1990), this study will tentatively consider the New Albany Shale as the source of these tars.

DISCUSSION

PETROLEUM CHARGE

Petroleum-charge values of 786 and 1,172 billion barrels derived from Rock-Eval expulsion factors are enormous quantities of petroleum. These quantities are one to two orders of magnitude greater than the quantity of discovered in-place petroleum within the basin, and it is difficult to rationalize these quantities as undiscovered petroleum or petroleum lost due to migration in or leakage from the basin. These unrealistic petroleum charge values can be attributed to Rock-Eval pyrolysis not simulating natural petroleum formation (i.e., generation and expulsion) and to the assumption that the fraction of inert organic carbon remains constant throughout petroleum formation.

High temperatures and low pressures used in Rock-Eval pyrolysis to vaporize generated hydrocarbons from a powdered rock do not occur naturally in subsiding sedimentary basins. In addition, the vaporized hydrocarbons from Rock-Eval pyrolysis represent a collective product with no distinction between high-molecular-weight hydrocarbons that are likely in nature to remain within the rock as bitumen, medium-to-low-molecular-weight hydrocarbons that are likely in nature to be expelled in a liquid phase as petroleum, and low-molecular-weight hydrocarbons that are likely in nature to be released as natural gas. This lack of distinction among natural phases broadens the predicted natural conditions for petroleum formation and exaggerates petroleum expulsion factors (Lewan 1993, 1994). Thermal decomposition of sedimentary organic matter may be discussed in terms of two opposing reaction pathways (Lewan, 1994). One pathway involves the cleavage of carbon-carbon bonds to yield medium-to-low-molecular-weight products (i.e., oil and gas). The other pathway involves the cross linking of carbon atoms to yield high-molecular-weight non-volatile products (i.e., pyrobitumen or inert carbon).

Rock-Eval pyrolysis promotes the cracking pathway and suppresses the cross-linking pathway by vaporizing and removing generated products from reaction sites within a rock before carbon-carbon cross-linking can occur. Similarly, this suppression of the cross-linking pathway is well demonstrated by oil-shale retorting methods, such as Fischer assays, which employ similar heating rates and temperatures as those used in Rock-Eval pyrolysis. Although Fischer assays use a cold trap to remove vaporized products instead of a carrier gas, their yields show a one-to-one correlation with Rock-Eval pyrolysis yields (Madec and Espitalié,

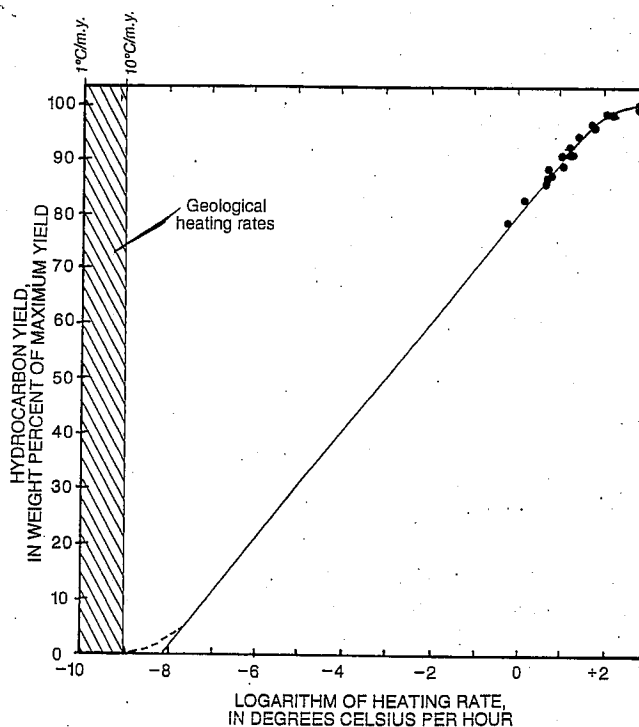


Figure 13. Relationship between percent hydrocarbon yield and heating rate of Fischer assays (Evans and Campbell, 1979) and extrapolation toward geological heating rates for sedimentary basins (i.e., 1°C/m.y. to 10°C/m.y.).

1984). The ability of these open anhydrous-pyrolysis methods to suppress the cross-linking pathway becomes apparent by examining the product yields at different heating rates. Evans and Campbell (1979) conducted Fischer assays over the same temperature range at different heating rates on aliquots of the same sample. Their results show product yields significantly decrease with decreasing heating rates (fig. 13). This relationship can be explained by slower vaporization and removal of generated products from reaction sites within the rock at lower heating rates, which results in promotion of the cross-linking pathway and reduced product yields. Extrapolation of this trend in figure 13 to lower heating rates results in no product yield at heating rates typically observed in subsiding sedimentary basins (i.e., 1°C/m.y. to 10°C/m.y.). The enormous quantities of expelled petroleum determined by Rock-Eval pyrolysis indicate the currently employed heating rates (25°C/min or 30°C/min) are too high, but, based on figure 13, it remains uncertain what heating rate is appropriate to simulate the balance between cracking and cross-linking pathways operative in natural petroleum formation.

Burnham (1991) has also observed reduced product yields with decreasing heating rates in open anhydrous pyrolysis and has questioned the assumption by Cooles and others (1986) that the fraction of inert organic carbon as

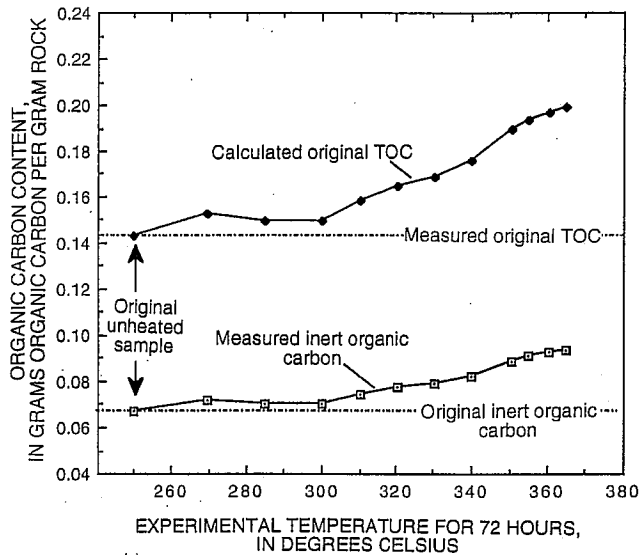


Figure 14. Change in measured inert organic carbon and calculated original total organic carbon (TOC^{om}) from Rock-Eval pyrolysis (Cooles and others, 1986) with increasing maturity, as induced by hydrous pyrolysis on aliquots of an immature sample of Clegg Creek Member of the New Albany Shale (fig. 11).

determined by Rock-Eval pyrolysis remains constant throughout petroleum formation. The rock recovered from the series of hydrous-pyrolysis experiments conducted on aliquots of shale from the Clegg Creek Member (fig. 11) provide samples at various stages of petroleum formation to further test this assumption. Rock-Eval data on the rock recovered from these experiments is given in table 9. The inert and original total organic carbon contents (C_1^0 and TOC^0) of the unheated rock are 67.5 and 143.4 mg/g of rock, respectively. These values give an inert organic carbon fraction of 0.471 (eq. 1), which is used with the data in table 9 to calculate the inert (C_1) and original total organic carbon contents (TOC^{om}) at different maturities induced by hydrous pyrolysis. Calculations are made with equation 2, which assumes that the inert organic carbon fraction (C_1) remains constant with increasing thermal maturity. The plotted results in figure 14 indicate that this assumption is not valid and that the amount of inert organic carbon, as determined by Rock-Eval pyrolysis, increases significantly with increasing thermal maturity. This increase results in significant overestimates of original total organic carbon (TOC^{om}) by as much as 39 percent (fig. 14). These overestimates of original total organic carbon in thermally mature rocks result in higher calculated organic carbon losses attributed to petroleum expulsion, which explain in part the unrealistic expulsion factors and enormous quantities of expelled petroleum derived from Rock-Eval pyrolysis.

These results suggest that, in addition to heating rate, the maturity level of a source rock can also influence the results from open anhydrous pyrolysis. This influence is not surprising when one considers the changes that occur to

organic matter and its distribution within a source rock with increasing thermal maturity. Sedimentary organic matter in thermally immature source rocks occurs predominantly as an insoluble solid kerogen, which is distributed within the mineral matrix as insular domains (Lewan, 1987). In naturally and experimentally matured source rocks, the kerogen domains decompose into a viscous, high-molecular-weight bitumen that impregnates the microporosity and parting planes of the rock to form a continuous organic network (Lewan, 1987). These state and distribution changes of the organic matter with increasing thermal maturity are also accompanied by compositional changes in the remaining bitumen as saturate-enriched petroleum is expelled from the rock (Lewan, 1993). Therefore, it is conceivable, with all of these physical and chemical changes occurring to the organic matter with increasing maturation, that the response from Rock-Eval pyrolysis would be different for the same rock at different maturity levels. The observed increase of inert organic carbon with increasing thermal maturation as shown in figure 14 indicates an increasing preference for the cross-linking pathway with increasing maturity of a source rock.

Although the petroleum charge determined by Rock-Eval pyrolysis is approximately 2.5 times greater than the petroleum charge determined by hydrous pyrolysis, the latter is still large relative to the discovered in-place petroleum within the basin. Hydrous pyrolysis exaggerates the natural processes responsible for petroleum expulsion (Lewan, 1994) and, therefore, should not be directly applied to sedimentary basins (Lewan, 1994). Three likely causes of this exaggeration by hydrous pyrolysis include enhanced solubility of H_2O in the bitumen network, reduced hydrostatic pressure, and enhanced thermal expansion of petroleum products.

Comparison of yields from closed pyrolysis experiments under hydrous and anhydrous conditions show the presence of water significantly suppresses the cross-linking pathway and promotes generation of soluble organic products (Comet and others, 1986; Tannenbaum and others, 1986; Lewan, 1993). Lewan (1992) suggests that the enhancement of yield under hydrous conditions relative to anhydrous closed conditions is a result of dissolved H_2O in the bitumen network acting as a source of hydrogen. It has been experimentally established that the solubility of water in hydrocarbons is two orders of magnitude higher than the solubility of hydrocarbons in water (Tsonopoulos and Wilson, 1983; Heidman and others, 1985; Skripka and Boksha, 1976; Guerrant, 1964; Griswold and Kasch, 1942). However, this high solubility of water in hydrocarbons is strongly dependent on temperature, with water solubilities being more than one order of magnitude higher at temperatures used in hydrous pyrolysis (300°C to 365°C) than at temperatures encountered in subsiding sedimentary basins (100°C to 200°C). A reduction of a possible hydrogen source by this magnitude is expected to promote the cross-linking pathway and reduce petroleum yields and expulsion factors under

Table 9. Rock-Eval data and calculated inert carbon and original total organic carbon in rocks recovered from hydrous-pyrolysis experiments on aliquots of a sample from the Clegg Creek Member of the New Albany Shale.[See fig. 11. See text for definitions of S₁ and S₂. HC, hydrocarbons; Rx, rock; wt.%, weight percent; TOC, total organic carbon; C, carbon]

Hydrous-pyrolysis conditions (°C for 72 h)	S ₁ HC (mg/g Rx)	S ₂ HC (mg/g Rx)	TOC (wt.% Rx)	Hydrogen index (S ₂ /g TOC)	C ₁ S ₁ -carbon (g C/g Rx)	C ₂ S ₂ -carbon (g C/g Rx)	Inert carbon (g C/g Rx)	Original TOC ¹ (g C/g Rx)
unheated	2.72	86.56	14.34	603.6	0.0023	0.0736	0.0675	0.1434
269.7	1.87	85.01	14.56	583.9	0.0016	0.0723	0.0718	0.1524
285.0	2.42	82.93	14.31	579.5	0.0021	0.0705	0.0706	0.1499
300.2	4.53	78.28	14.09	555.6	0.0039	0.0665	0.0705	0.1497
310.1	6.39	69.56	13.90	500.4	0.0054	0.0591	0.0744	0.1580
320.3	8.24	57.05	13.31	428.6	0.0070	0.0485	0.0776	0.1648
330.1	11.49	44.37	12.70	349.4	0.0098	0.0377	0.0795	0.1688
340.2	10.67	33.71	12.04	280.0	0.0098	0.0287	0.0827	0.1756
350.5	7.44	19.56	11.21	174.5	0.0063	0.0166	0.0892	0.1894
355.1	6.37	12.50	10.72	116.6	0.0054	0.0106	0.0912	0.1936
360.3	5.05	10.94	10.64	102.8	0.0043	0.0093	0.0928	0.1970
365.0	4.70	9.32	10.58	88.1	0.0040	0.0079	0.0939	0.1994

¹Inert organic carbon fraction = 0.471 (see eq. 1).

natural conditions. Research is currently under way to determine the degree to which hydrous pyrolysis exaggerates expulsion in sedimentary basins (Lewan, 1994), but the differences in expulsion between anhydrous and hydrous pyrolysis (Lewan, 1993) suggest that exaggeration by a factor of 2 is conceivable.

Several experimental studies have demonstrated that increasing hydrostatic pressure can significantly reduce hydrocarbon yields from source rocks (Price and Wenger, 1992; Freund and others, 1992, figs. 2 and 3; Michels and others, 1992). Hydrous pyrolysis typically operates at hydrostatic pressures slightly higher than the vapor pressure of water as determined by the experimental temperature and amount of gas generated from the pyrolyzed rock (Lewan, 1993). The hydrous-pyrolysis experiments conducted at 330°C for 72 hours by Michels and others (1992) showed a total yield (bitumen + expelled oil) reduction of 34 weight percent at a pressure of 70 MPa and 58 weight percent at a pressure of 130 MPa relative to the total yields under typical hydrous-pyrolysis conditions at a pressure of 13 MPa. According to paleoreconstructions of the Illinois Basin, the New Albany Shale was within peak petroleum generation at depths from 2.44 to 3.36 km (Cluff and Byrnes, 1990). If the New Albany Shale at these depths was under normal hydrostatic conditions (10.2 kPa/m), the water pressure would be 24.9 to 34.3 MPa. Under overpressure conditions, approaching lithostatic pressures (20.4 kPa/m), the water pressure at these same depths would be 49.8 to 68.5 MPa (Hamilton-Smith, 1993). This latter range of water pressures would correspond approximately to a 25 to 35 percent reduction in yield based on the hydrous-pyrolysis results of Michels and others (1992) at 330°C for 72 hours.

The higher temperatures used in hydrous pyrolysis relative to natural petroleum generation may also exaggerate expulsion factors because of greater thermal expansion of generated petroleum within source rocks. The specific volume of petroleum with a 29° API gravity is 1.21 cm³/g at 100°C and 1.51 cm³/g at 315°C (GPSA, 1987). Therefore, a given mass of petroleum generated in a source rock by hydrous pyrolysis at 315°C will occupy 25 percent more volume than that generated in a source rock by natural subsidence to 100°C. This percent of volume increase may reflect an equivalent increase in petroleum expulsion during hydrous pyrolysis. Most petroleum expulsion models require the porosity of a source rock to be filled with bitumen or petroleum before expulsion occurs (e.g., Durand and others, 1987; Leythaeuser and others, 1987). Lewan (1987) reports notable expansion of source-rock cores recovered from hydrous-pyrolysis experiments as a result of the opening of oil-filled parting fractures parallel to the bedding fabric. Although petroleum densities at temperatures in excess of 320°C are not available, extrapolation of the lower temperature trends (<320°C) suggests volume increases may exceed 30 percent at experimental temperatures of 350°C.

Although hydrous pyrolysis simulates natural petroleum formation, as indicated by the composition of its expelled oil and petrographic character of its recovered rock, its expulsion factors appear to be exaggerated to some degree compared to natural expulsion factors. Additional research is critical to determining the degree of exaggeration. However, the general evaluation of water solubility, hydrostatic pressure, and thermal expansion effects suggest that it would be reasonable to assume natural expulsion factors to be at least one-half of the hydrous-pyrolysis expulsion factors. This temporary estimated correction to the

hydrous-pyrolysis expulsion factors would reduce the petroleum charge values to 152 and 236 billion barrels for the dilution and steady-state models, respectively.

PETROLEUM LOSSES

Petroleum losses within the Illinois Basin may occur as residual oil in porous-rock carrier beds; surface leakage in the form of tar seeps; and eroded traps, carrier beds, or seeps in the post-Pennsylvanian section. More than 90 percent of the accountable produced petroleum within the Illinois Basin occurs in reservoirs that are stratigraphically at least 350 m above the New Albany Shale. No continuous porous strata capable of serving as a regional carrier bed occurs within this intervening section between source and reservoir. As a result, it becomes necessary to advocate petroleum migration through regional fractures and faults. Secondary migration through fractures and faults has been advocated as effective pathways in other sedimentary basins (e.g., Bell, 1989; Price, 1980; du Rouchet, 1984). Evidence of this mode of secondary migration in the Illinois Basin includes the association of tar sands with faults in western Kentucky (Williams and others, 1982), lenticular reservoirs encased in impervious strata of the Ste. Genevieve Limestone and Cypress Sandstone (Seyler and Cluff, 1990), and the apparent migration of petroleum along faults to account for petroleum in Pennsylvanian reservoirs (Swann and Bell, 1958; Seyler and Cluff, 1990).

Although secondary migration through fractures and faults appears to be significant in the Illinois Basin, porous carrier beds may have been critical to lateral migration of petroleum away from faults and fracture conduits. The regional extent and good reservoir character of the Aux Vases Sandstone (Mississippian) make it a strong candidate as a regional carrier bed. This possible carrier bed varies in thickness from an average of 7.6 m in the central part of the basin to 61 m in the western part of the basin (Seyler and Cluff, 1990), and it is an exploration target from which almost 1 billion barrels of petroleum have been produced (Treworgy, 1990). It has a mean porosity of 18 percent and a mean permeability of 76 mD based on data from 179 fields (Mast, 1970). The Sylamore Member (Devonian), which occurs at the base of the New Albany Shale, has also been suggested as a possible carrier bed (Oltz and Crocket, 1989; Bethke and others, 1991). However, this rock unit varies in thickness from a few centimeters to 1.6 m and has a discontinuous distribution, as mapped by Workman and Gillette (1956), that make its significance as a regional carrier bed unlikely.

The importance of considering porous-rock units as carrier beds is the ability of granular rocks to retain residual petroleum after they have served as carrier beds (Schowalter, 1979). This residual petroleum in carrier beds has been

advocated as the cause of the enormous quantities of petroleum that must have been lost, based on the enormous quantities of expelled petroleum determined by Rock-Eval pyrolysis (Mackenzie and Quigley, 1988). Schowalter (1979) suggests that only the upper few feet of a potential carrier bed experiences petroleum migration, and England and others (1987) suggest that only 1 to 10 percent of the cross-sectional area of a carrier bed experiences petroleum migration. The amount of residual petroleum that remains in this portion of carrier bed that experiences petroleum migration ranges from 10 to 30 percent of the pores for water-wet sandstones (Schowalter and Hess, 1982). The surface area of the Aux Vases Sandstone that contains a sandstone-to-shale ratio greater than 1:8, as mapped by Swann and Bell (1958), is 54.5×10^9 m². Assuming that 3 m of this entire area acted as a carrier bed with a mean porosity of 18 percent, 10 to 30 percent of this pore volume equates to 18.5 to 55.5 billion barrels of unrecoverable residual petroleum lost in the carrier bed. Currently, these quantities cannot be substantiated by petroleum shows from drilling reports or cores. Therefore, this range is treated as a maximum loss that includes residual petroleum that may have occurred on a more local scale in other sandstones younger than the Aux Vases Sandstone.

Petroleum may also be lost from a basin as a result of untrapped petroleum leaking out of carrier beds, faults, or fractures at or near the surface. Although weathering of leaking petroleum by evaporation, water washing, microbial degradation, polymerization, and auto-oxidation (Hunt, 1979, p. 411) may result in as much as 90 percent loss of the original petroleum, a residual tar will remain to indicate the occurrence of petroleum leakage. The only major tar deposits within the Illinois Basin occur in sandstones that subcrop at or immediately overlie the Mississippian-Pennsylvanian unconformity in western Kentucky. These tar-sand deposits comprise an in-place resource of 3.4 billion barrels, with 1.73 billion barrels being measured and 1.68 billion barrels being speculative (Kuuskraa, 1984; Noger, 1987; Crysdale and Schenk, 1988). The API gravity of the tar is typically 10° (Ward, 1983; Fedde and Zavada, 1983; Moore and others, 1984), with a mean sulfur content of 1.66 weight percent (Ward, 1983; Fedde and Zavada, 1983; Hosterman and others, 1990; Thomas and others, 1983). These properties are significantly different than those of the produced petroleum of the Illinois Basin, which has a mean API gravity of 36.8° (Mast, 1970) and a mean sulfur content 0.22 weight percent (Huff, 1987).

Subsurface degradation of produced petroleum to tar with these properties has been documented in the Brea Olinda field of the Los Angeles Basin (W.L. Orr, written commun., 1994). Microbial activity or water washing is the usual cause of this degradation (Bailey and others, 1973). These processes do not appear to add or remove sulfur, which occurs predominantly in the more aromatic and high-molecular-weight fraction of most petroleum (W.L. Orr, written commun., 1994). As a result, sulfur content can be

used to determine the amount of petroleum lost from the tar-sand deposits prior to degradation. Mass-balance calculation using the mean sulfur content of 0.22 weight percent for the original petroleum from which the tar was derived indicates that the 3.4 billion barrels of in-place tar represents only 13.3 weight percent of the original in-place petroleum. This percentage is in agreement with the study by Tannenbaum and others (1987), which determined that the tars in the Dead Sea area represented 10 to 20 percent of the original undegraded petroleum. Using 13.3 percent and correcting for density changes from 36.8° to 10° API gravity gives 30.4 billion barrels of original in-place petroleum prior to degradation. Approximately half of this total is measurable barrels (15.4×10^9 bbl) and the remaining half is speculative barrels (15.0×10^9 bbl). This total quantity is more than 2.5 times greater than the 12 billion barrels of discovered in-place petroleum and represents 27 billion barrels of petroleum lost from the basin by degradation of near-surface leakage.

The final petroleum loss to evaluate in the Illinois Basin is the amount of petroleum originally hosted in the eroded Upper Pennsylvanian through Permian section. Thermal-history models based on vitrinite reflectance and moisture content of Lower Pennsylvanian coal indicate that 0.3 to more than 2 km of Upper Pennsylvanian through Permian section was removed by erosion during the Mesozoic and Cenozoic (Damberger, 1971; Cluff and Byrnes, 1990). Petroleum formation from the New Albany Shale was well underway by the end of the Permian (Cluff and Byrnes, 1990), which makes it possible for petroleum to have migrated into carrier beds, traps, and seeps within the Upper Pennsylvanian through Permian section. The likelihood of this possibility is supported by surface tar deposits in Lower Pennsylvanian sandstones in western Kentucky and by shallow petroleum production from Lower Pennsylvanian sandstones at subsurface depths between 50 and 150 m in Illinois and Indiana (Ball Associates, 1965). However, the amount of petroleum that may have migrated into the Upper Pennsylvanian through Permian section remains uncertain.

Petroleum production from Lower Pennsylvanian sandstones comprise 13 percent of the total basin production, with 88 percent of it occurring in traps associated with the LaSalle anticline (Seyler and Cluff, 1990). Currently, field studies are not available on whether these traps are filled to their spill points. If these traps are not filled to their spill points and have seals capable of maintaining petroleum columns to the spill points, it may be assumed that the petroleum charge was less than spill-point volumes and migration of petroleum into the overlying Upper Pennsylvanian through Permian section was minimal. Similar field studies on Mississippian traps associated with the Salem, Loudon, and Johnsonville anticlines would also be helpful in evaluating whether a limited petroleum charge in these deeper parts of the basin prevented petroleum accumulations from

occurring in overlying closures in Lower Pennsylvanian sandstones. Until field studies of this type are made within the context of migration pathways and petroleum charge in Illinois and Indiana, no defensible estimate can be made on the quantity of petroleum lost due to erosion of the Upper Pennsylvanian through Permian section.

The petroleum losses associated with the eroded section overlying the tar deposits of western Kentucky is a more facile consideration. The 3.4 billion barrels of tar that equate to 30.4 billion barrels of original undegraded petroleum occurs at subsurface depths shallower than 200 m (Noger, 1987). Cluff and Byrnes (1990) estimate that more than 600 m of section have been removed from this area by erosion. These tar deposits are very different in size and geometry from oil reservoirs and appear to be truncated migration pathways that may have fed reservoirs in the Upper Pennsylvanian through Permian section prior to erosion. If it is assumed that the areal extent of petroleum removed by erosion was the same as that remaining, the original volume of undegraded petroleum removed by erosion would be 91.2 billion barrels. Approximately half of this total is derived from the measurable barrels (46.2×10^9 bbl), and the remaining half is derived from the speculative barrels (45.0×10^9 bbl). The sum of this lost petroleum with the petroleum of the remaining tar deposits (i.e., 30.4 billion bbl) is 121.6 billion barrels of undegraded petroleum, which equates to 13.6 billion barrels of tar prior to erosion. Although this quantity appears large relative to the 12 billion barrels of in-place petroleum discovered in the Illinois Basin, it is not excessive compared to the 12.5 to 16 billion barrels of tar in the Tar Sand Triangle of Utah (Crysdale and Schenk, 1988) or the 626 billion barrels of heavy bitumen in the Athabasca tar deposit of Canada (Carrigy, 1974).

MATERIAL BALANCE

The material balance between accountable petroleum in the Illinois Basin and petroleum charge from the New Albany Shale is given in table 10. The petroleum charge based on one-half of the expelled oil from hydrous-pyrolysis yields the most reasonable values of 152 to 236 billion barrels, compared with the 92 to 189 billion barrels of accountable petroleum in the basin. Approximately two-thirds of this accountable petroleum may have migrated along fault systems and the Mississippian-Pennsylvanian unconformity in western Kentucky—only 6 to 13 percent migrated into discovered reservoir traps. The remaining 18.5 to 55.5 billion barrels of petroleum may have been lost as residual petroleum along migration pathways in porous carrier beds.

Although this material balance yields reasonable values, the uncertainty in assumptions and variability in values make it too equivocal to determine the amount of petroleum remaining to be discovered. The undiscovered quantities of petroleum sourced from the New Albany Shale in the Illinois

Table 10. Material balance of accountable petroleum and petroleum charge from source rocks of the New Albany Shale in the Illinois Basin.

Billions of barrels	
Accountable petroleum	
In-place discovered	12.0
Residual in carrier beds	18.5-55.5
Undegraded tar deposits	15.4-30.4
Lost by erosion	<u>46.2-91.2</u>
Total accountable petroleum	92.1-189.1
Petroleum charge	
½ hydrous pyrolysis	152-236*
Hydrous pyrolysis	305-472*
Rock-Eval pyrolysis	786-1,172*
Undiscovered petroleum **	
Least difference	-37
Greatest difference	144

*Range based on dilution and steady-state models for organic richness in source rocks.

**Based on ½ hydrous-pyrolysis expulsion factor.

Basin show a wide range of values. Using the one-half hydrous-pyrolysis expulsion factors, a minimum of -37 billion barrels of undiscovered petroleum is obtained by subtracting the maximum accountable petroleum value (189.1) from the minimum petroleum charge value (152). This deficit may be interpreted as indicating that no more petroleum from the New Albany Shale remains to be found in the Illinois Basin or that estimates of erosional or residual losses are too high. Conversely, a maximum of 144 billion barrels of undiscovered petroleum from the New Albany Shale may be obtained from the same expulsion factors by subtracting the minimum accountable petroleum value (92.1) from the maximum petroleum charge value (236). However, the intensity of drilling in the Illinois Basin makes this maximum unlikely. Mast and others (1989) estimate that only 0.46 billion barrels of undiscovered petroleum remains to be produced in this Illinois Basin, and 75 percent of this petroleum will occur in fields smaller than 1 million barrels (Oltz and others, 1990). This estimated amount of undiscovered petroleum equates to 1.3 billion barrels of in-place petroleum by using a recovery factor of 35 percent. The actual amount of remaining undiscovered petroleum expelled from the New Albany Shale is considered to be somewhere between these two extremes of 0 and 144 billion barrels.

The question that emerges from this wide range of undiscovered petroleum is whether or not a more accurate material-balance assessment can be achieved through collection of new data and further research. An accurate petroleum charge is obviously one of the critical concerns. The amount of accountable petroleum will always be subject to differing interpretations due to uncertainties in evaluating migration and erosional losses, but realistic and well-constrained

petroleum-charge values can establish an ultimate quantity of petroleum expelled from mature and over-mature source rocks within a basin. This realistic ultimate petroleum charge for a basin can serve as a fixed quantity from which interpretations can be evaluated, compared, and discussed. Establishing this ultimate quantity of petroleum requires accurate characterization of the source rock and realistic expulsion factors. The importance of an accurate characterization of a source rock is illustrated by the 84 billion barrel difference between the dilution and steady-state models using the same one-half hydrous-pyrolysis expulsion factors (table 10). Similarly, the importance of realistic expulsion factors is demonstrated by the 481 to 700 billion barrel difference between petroleum charges derived from Rock-Eval and hydrous-pyrolysis expulsion factors (table 10). In addition to developing a thorough understanding of the factors that influence laboratory pyrolysis yields, it is also essential to know how they relate to natural maturity intervals within a sedimentary basin. This feasibility study shows that 45 to 50 percent of the petroleum charge derived from Rock-Eval expulsion factors and 34 to 39 percent of the petroleum charge derived from the hydrous-pyrolysis expulsion factors are from source rocks in the 0.6- to 0.7-percent R_o interval (tables 7 and 8). If further studies reveal that this early stage of maturation does not yield these quantities under natural conditions, the resulting petroleum-charge values may be significantly reduced by these percentages. As an example, the petroleum charge using the one-half hydrous-pyrolysis expulsion factors for maturity interval greater than 0.7 percent R_o is 93 to 157 billion barrels. Clearly, accurate expulsion factors and source-rock characterization are essential to the utility of material-balance assessments.

FUTURE RESEARCH

This material-balance assessment relied solely on suppressed vitrinite reflectance in the New Albany Shale to determine level of thermal maturity, which was correlated to Rock-Eval hydrogen indices to determine degree of petroleum generation. Other parameters related more specifically to petroleum expulsion from a source rock need to be researched. Lewan (1985) demonstrated that type-II kerogens isolated from source rocks subjected to hydrous pyrolysis had a common relationship between their atomic H/C ratio and percent of expelled oil, irrespective of differences in their kinetics of petroleum formation. Stages of petroleum generation and the effectiveness of potential source rocks have also been evaluated by petrophysical log properties (Meissner, 1978) and petrographic studies (Lewan, 1987). Preferential expulsion of particular classes of biomarker compounds (e.g., monoaromatic steroid hydrocarbons) from source rocks (Lewan and others, 1986) may also provide an index of petroleum expulsion. Focused research efforts on these and other possible indicators of petroleum generation

and expulsion are attainable and relevant to obtaining more accurate material-balance assessments in the future.

Additional research is needed to understand the natural factors that control petroleum expulsion from source rocks. Hydrous-pyrolysis experiments have shown that the presence of water (Lewan, 1993), hydrostatic pressure (Michels and others, 1992), mineral matrix (Lewan and Whitney, 1993), and amount of organic matter (Lewan, 1987) can have a significant effect on petroleum expulsion. These factors need to be quantified and related to chemical and physical mechanisms responsible for petroleum generation and expulsion for reasonable extrapolations to natural conditions. Specific to this material balance on the Illinois Basin, additional experimental work is needed on other members of the New Albany Shale. This study assumed that the experimental results on the sample from the Clegg Creek Member are representative of source rocks in all of the members of the New Albany Shale.

Collection of source rocks and examination of their relationship to carrier beds in the subsurface also needs further study. Price (1994) maintains that the petroleum potential remaining in a mature source rock is underestimated because of significant petroleum losses to the drilling mud as cuttings or cores move up the borehole for collection at the wellhead. If this phenomenon is common, expulsion factors will be overestimated and a major petroleum resource remaining within source rocks may be unrecognized. In addition, the thickness of a source rock and its relationship to adjacent carrier beds may significantly influence expulsion factors. Leythaeuser and others (1988) have presented geochemical data on cores from the North Sea that suggest expulsion efficiencies are greater at the margins of thick source rocks where they are in contact with carrier beds. This interpretation can be tested in the Illinois Basin at the contact between the Sylamore Member and overlying organic-rich shales in the New Albany Shale.

Although an accurate petroleum charge is critical to establishing an upper limit for a material-balance assessment, an accurate determination of petroleum losses during secondary migration is critical to establish the amount of undiscovered petroleum remaining within a basin. Detailed studies on migration pathways of petroleum from the mature source rock to convergence in traps or divergence throughout the basin are needed. Petroleum accumulations and seeps along migration pathways need to be evaluated in terms of spill points, seal integrity, and related accumulations. Secondary migration through regional or tectonic fractures, fault zones, or permeable porous strata need to be evaluated from buoyancy and fluid-flow models, compositional data on fluid inclusions, secondary porosity occurrences and timing, petroleum shows and staining, and compositional changes in migrated petroleum. A major assumption in this feasibility study is that tar deposits in western Kentucky were derived from petroleum expelled from the New Albany Shale. Although organic geochemical studies have correlated the

petroleum in Illinois with the New Albany Shale (Hatch and others, 1990), no geochemical correlations have established the source of petroleum and tar in western Kentucky. These tars may account for more than two-thirds of the accountable petroleum within the basin. Therefore, an organic geochemical study to establish the source rock of these tars is essential to future material-balance assessments of the Illinois Basin.

Although this future research represents a major scientific effort, the focusing of this basic research effort has immediate applications to providing more accurate and credible petroleum resource assessments.

CONCLUSIONS

Critical to a material-balance assessment is the petroleum charge of source rocks within a basin. The two essential components for an accurate petroleum charge are well-characterized source rocks and realistic petroleum expulsion factors. Although the thermally immature cores and outcrops of the New Albany Shale are well characterized, the effect of significant thickening of this rock unit on its organic richness at higher maturity levels has not been well characterized. The steady-state and dilution models used in this study serve as good end members in projecting organic richness of the source rocks in their depocenter, but more accurate and member-specific models for organic richness are needed. This enhanced accuracy is attainable through better characterization of source rocks in cores representing higher levels of thermal maturity within the basin.

It also becomes apparent from this study that results from laboratory pyrolysis methods cannot be directly used to determine expulsion factors for naturally matured source rocks. The unnatural method by which a product is derived in Rock-Eval pyrolysis and the incorrect assumption that inert carbon of a source rock remains constant with increasing thermal maturity results in unrealistic petroleum-charge values ranging from 786 to 1,172 billion barrels of petroleum. Although hydrous pyrolysis simulates the natural processes involved in petroleum generation and expulsion, the higher temperatures and lower hydrostatic pressures in these experiments compared with natural conditions results in exaggerated expulsion factors. More research is needed to evaluate the degree of exaggeration, but a qualitative evaluation suggests that hydrous pyrolysis may exaggerate expulsion factors by a factor of 2. Using this factor to correct expulsion efficiencies derived from hydrous pyrolysis gives more realistic petroleum-charge values ranging from 152 to 236 billion barrels of petroleum. Thirty-four to thirty-nine percent of this petroleum charge is derived from source rocks with thermal maturities between 0.6 and 0.7 percent R_o . If future studies indicate that petroleum expulsion did not occur within this maturity interval, then the petroleum charge for the basin could decrease to a range of 93 to 157 billion barrels.

Accounting for the amount of petroleum lost within or from a basin is also critical to a material-balance assessment. This aspect requires accurate determinations of residual petroleum in porous carrier beds, surface leakage because of insufficient traps along migration pathways, and removal of reservoired petroleum by erosion. Secondary migration in the Illinois Basin appears to have been dominated by vertical migration along fractures and faults. The significance of this style of migration is demonstrated by the fault-related tar sands of western Kentucky, which represent 15 to 30 billion barrels of undegraded petroleum. An additional 46 to 91 billion barrels of petroleum may have existed in overlying reservoirs before they were eroded during post-Permian time. These values suggest that more than two-thirds of the accountable petroleum expelled from the New Albany Shale migrated through fractures or faults into western Kentucky. This study also suggests that as much as 18 to 56 billion barrels of residual petroleum may reside in the Aux Vases Sandstone and other porous rocks, which may have acted in part as a porous carrier bed for lateral migration. It becomes apparent from this type of assessment that additional research on delineating migration pathways and understanding fluid flow from source rocks to traps is essential.

Petroleum material-balance assessments of established petroleum basins are feasible. Critical to this method is accurate determinations of petroleum charge from a source rock, discovered in-place petroleum, and petroleum losses due to migration, leakage, and erosion. Although this material balance yields reasonable values, the uncertainty in assumptions and variability in values make it too equivocal at this time to determine the precise amount of petroleum remaining to be discovered. This feasibility study yields a broad range of -37 to 144 billion barrels of petroleum from the New Albany Shale that remain to be discovered in conventional or unconventional accumulations of the Illinois Basin. The -37 billion barrels indicates that possibly no more undiscovered petroleum remains to be found or that the amounts of accountable petroleum loss to carrier beds or erosion are overstated. The 144 billion barrels indicates that possibly the expulsion factors or the steady-state model for organic richness for the source rocks in the mature part of the basin are overstated. Although this range is broad, it does encompass and does not completely exceed the range of accountable petroleum within the basin. As further data and research refine this type of assessment, this broad range of undiscovered petroleum is expected to narrow and provide a more precise upper limit to statistical-analog assessments currently employed. In addition, the quantitative understanding gained by material-balance assessments of established petroleum basins will provide a more scientific method for assessing petroleum in frontier basins on a global scale.

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